

# Chondritic meteorites and their components: Constraints on the early Solar System processes

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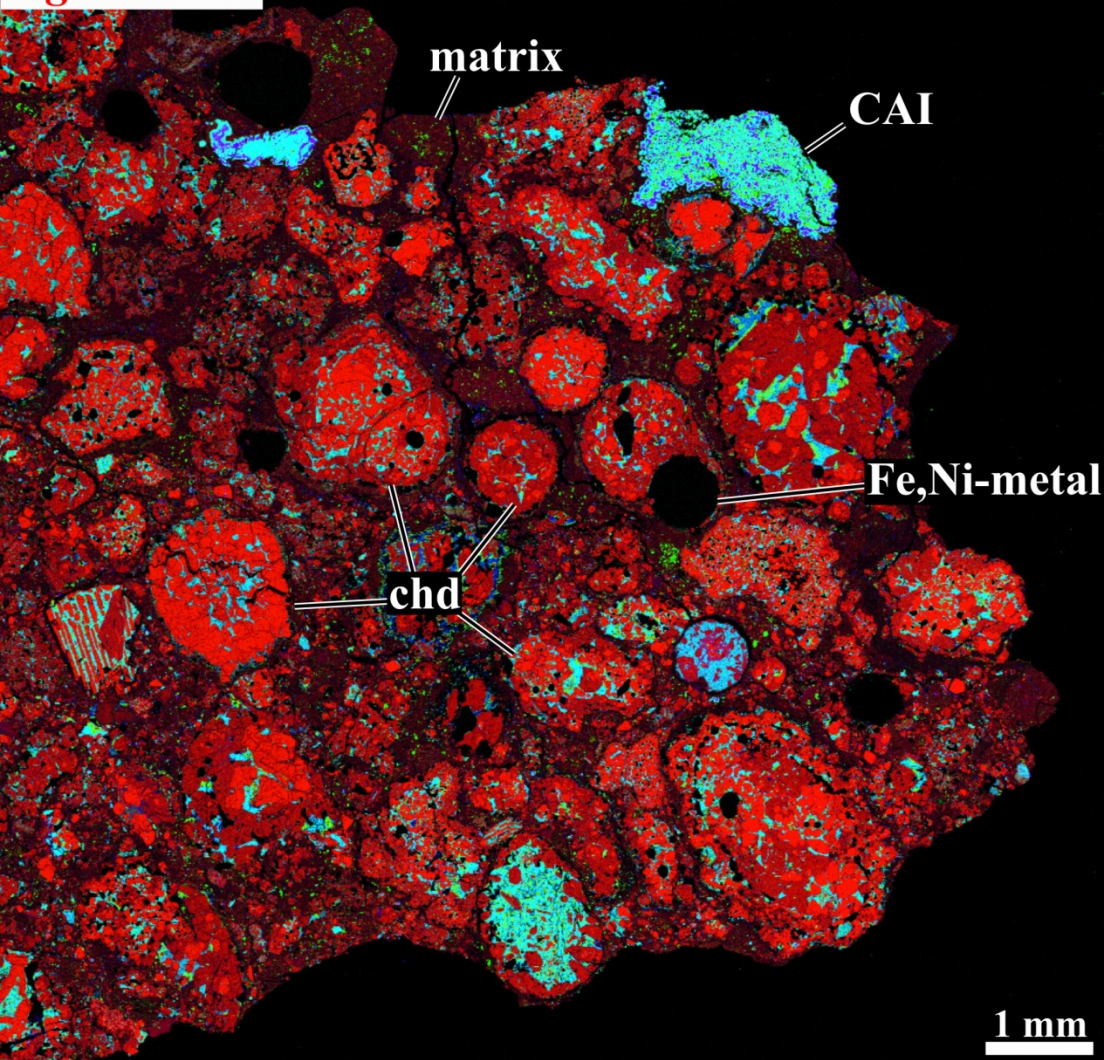
# Outline

- I. Introduction: Chondritic meteorites & their classification
- II. Calcium-Aluminum-rich Inclusions (CAIs)
  - constraint on astrophysical setting of Solar System formation
  - constraint on oxygen-isotope reservoirs
- III. Chondrules & fine-grained matrices
  - constraint on life-time of protoplanetary disk (PPD)
  - constraint on duration of accretion of chondrite parent asteroids
- IV. Aqueous activity on chondrite parent asteroids
  - constraint on sources of asteroidal water



# Major chondritic components: Chondrules, matrix, & CAIs

**Mg : Ca : Al**



CR carbonaceous chondrite • chondrules + Fe,Ni-metal

(30–98 vol%)

• Ca,Al-rich inclusions (CAIs)  
(<1–5 vol%)

• fine-grained matrix  
(<2–70 vol%)

**dominated by crystalline material**

→ thermal processing in PPD  
(evaporation, condensation, thermal  
annealing, & melting)



# Classification of chondritic meteorites (chondrites)

- based on mineralogy, petrography, bulk oxygen-isotope & chemical compositions, chondrites are divided into 15 groups & 3 major classes

<u>Carbonaceous</u>							<u>Enstatite</u>		<u>Ordinary</u>			Other		
CI	CM	CR	CV	CK	CO	CB	CH	EH	EL	H	L	LL	K	R

- letters designating groups refer to a prototype meteorite in a group:

CI – Ivuna-like (CI chondrite: Orgueil)

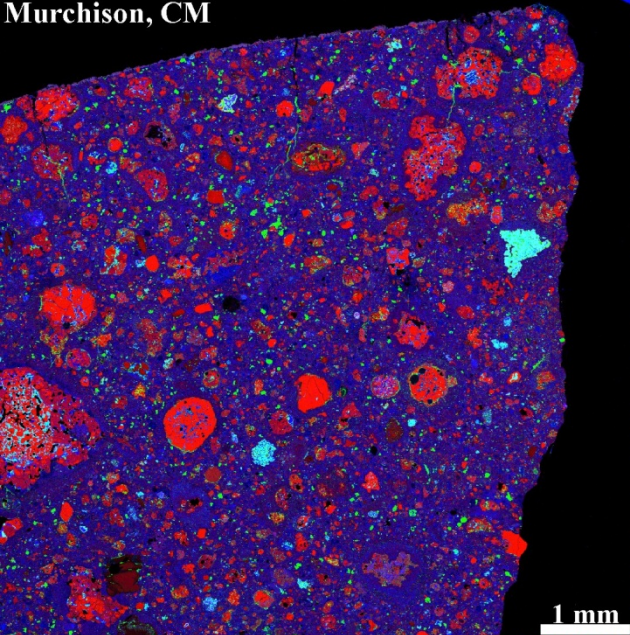
CM – Mighei-like (CM chondrite: Murchison)

CV – Vigarano-like (CV chondrite: Allende)

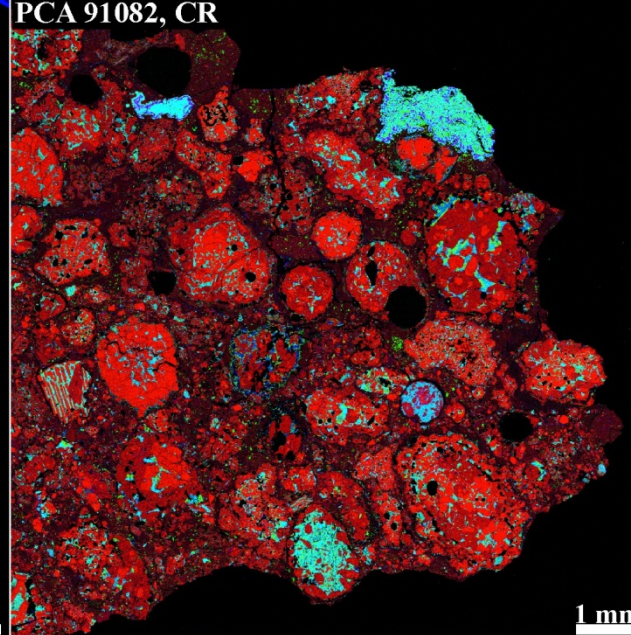
- some chondrites are ungrouped (e.g., Acfer 094)



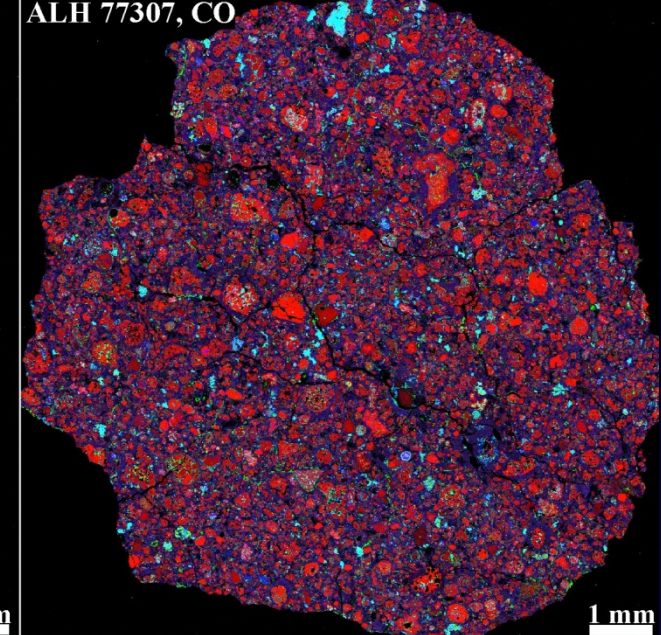
Murchison, CM



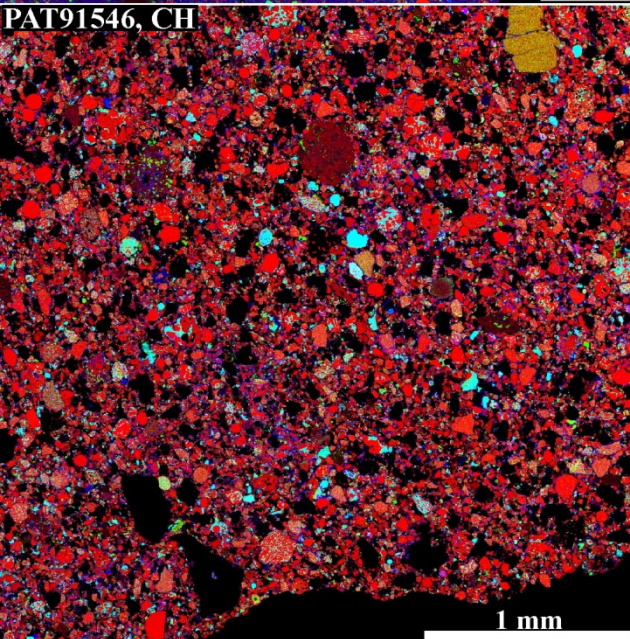
PCA 91082, CR



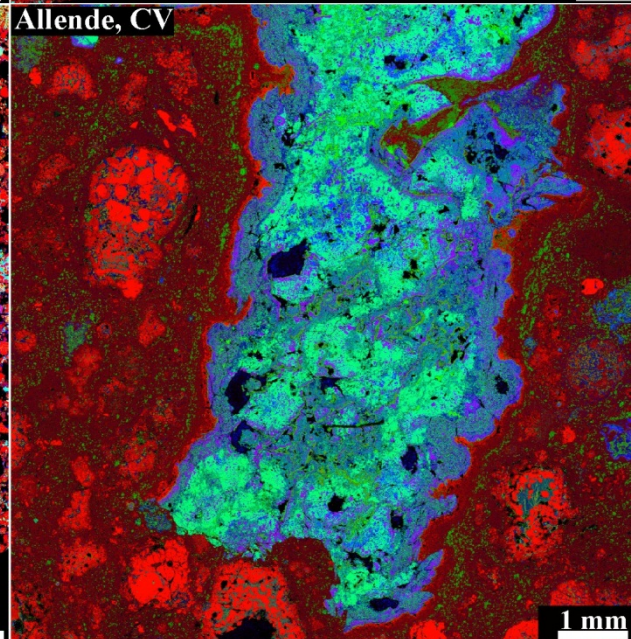
ALH 77307, CO



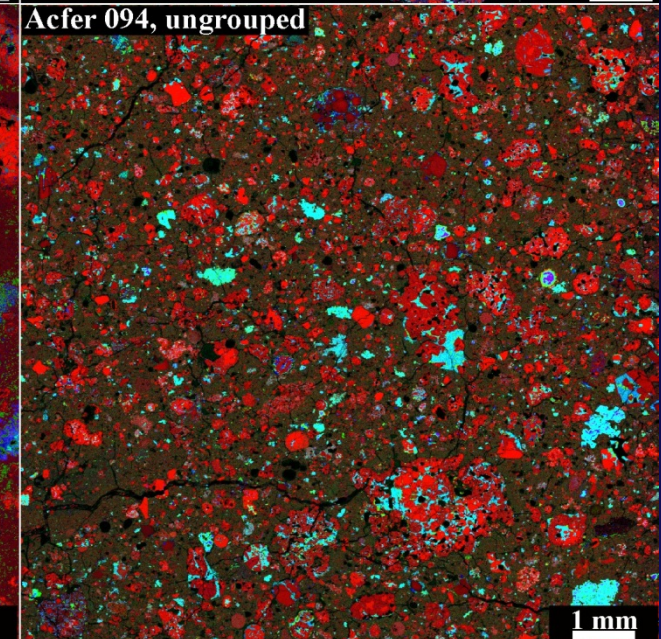
PAT91546, CH



Allende, CV

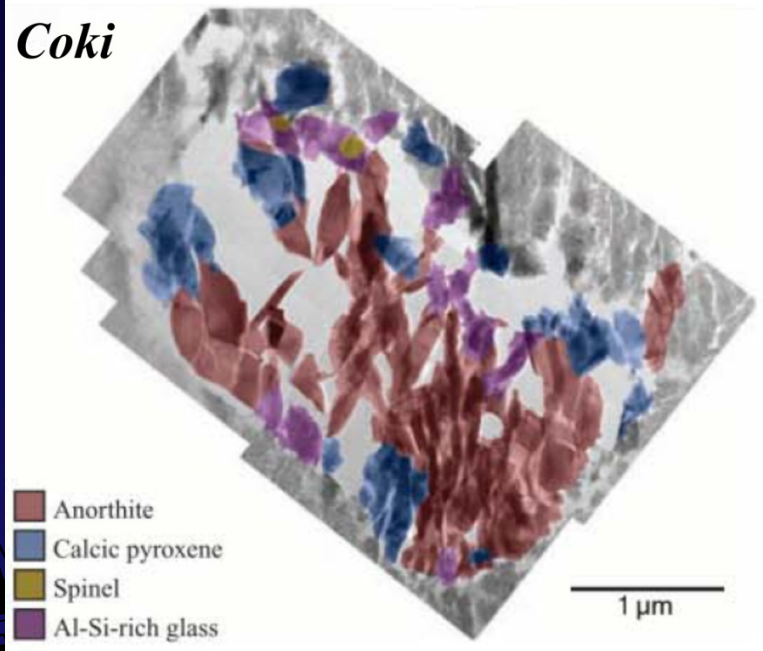
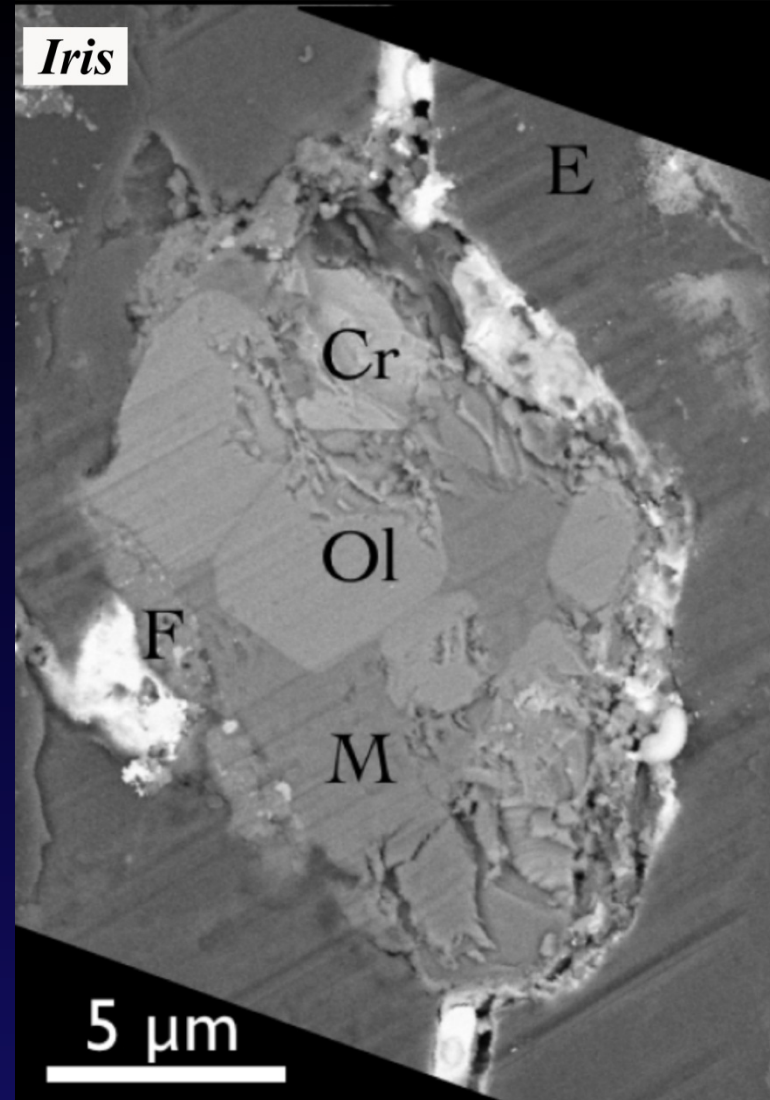
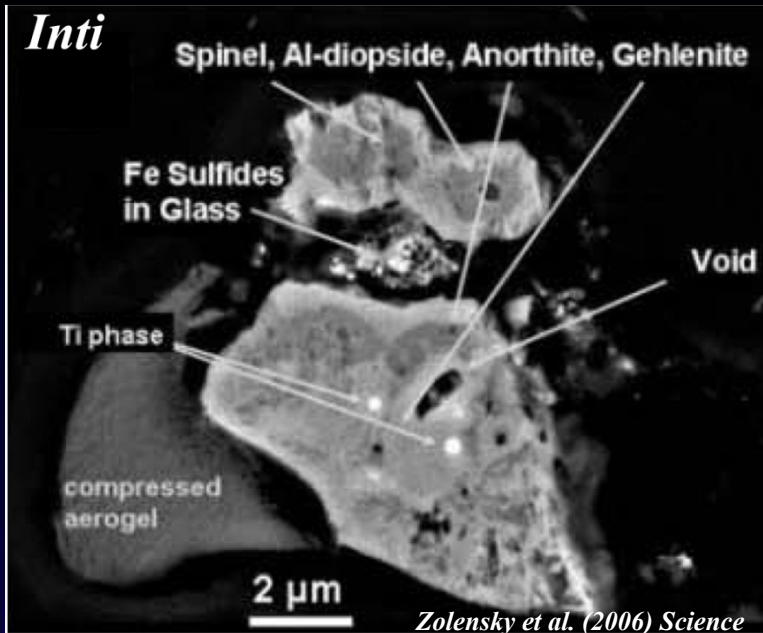


Acfer 094, ungrouped

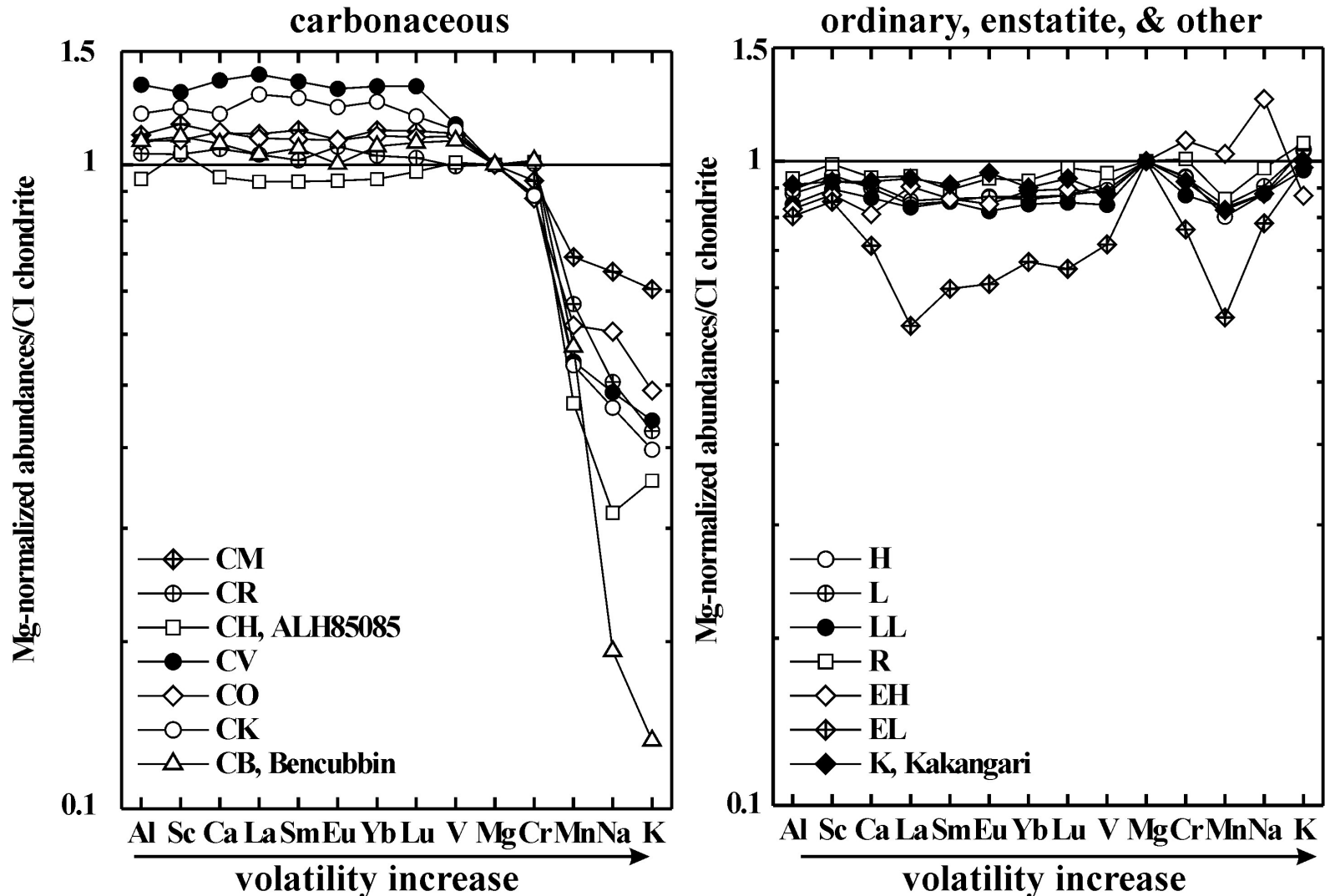


- intergroup variations in textures, mineralogy, sizes & abundances of chondritic components
- CAIs & chondrules present in all chondrite groups & in a comet *81P/Wild 2*

# CAIs & a chondrule fragment from 81P/Wild 2 comet



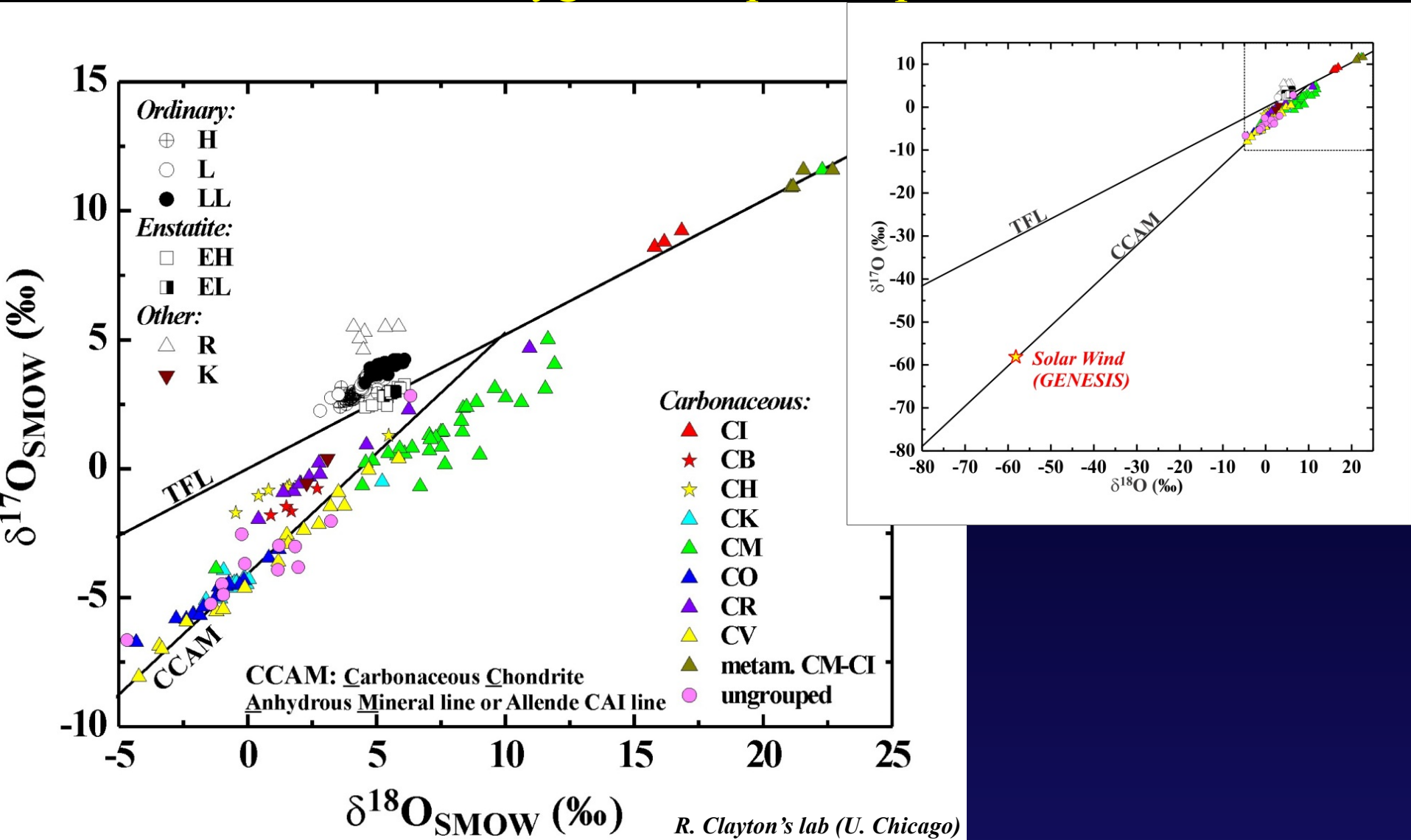
# Bulk chemical compositions



*J. Wasson's lab (UCLA)*

- CIs compositionally most similar to solar photosph. & in cosmochemistry CIs ≡ Sun
- chondrite groups have distinct bulk chemical compositions
- chondrites within a group are compositionally similar → sampled same asteroid

# Bulk oxygen-isotope compositions



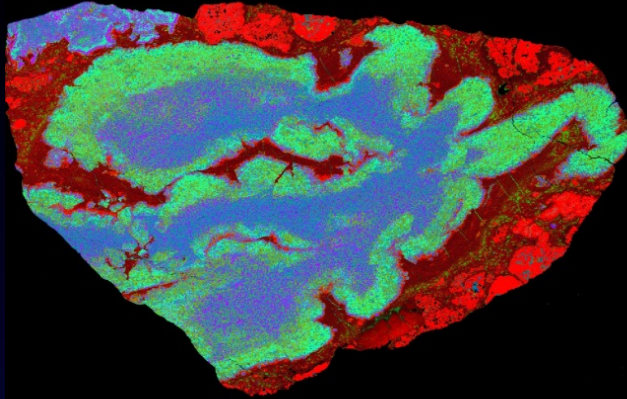
$$\delta^{17,18}\text{O} = \left[ \frac{(^{17,18}\text{O}/^{16}\text{O})_{\text{sample}}}{(^{17,18}\text{O}/^{16}\text{O})_{\text{SMOW}}} - 1 \right] \times 1000, \text{ Standard Mean Ocean Water}$$

- chondrite groups occupy distinct regions on three-isotope oxygen diagram
- very different from composition of solar wind,  $\delta^{17,18}\text{O} \approx -60\text{‰}$  (McKeegan et al., 2011)



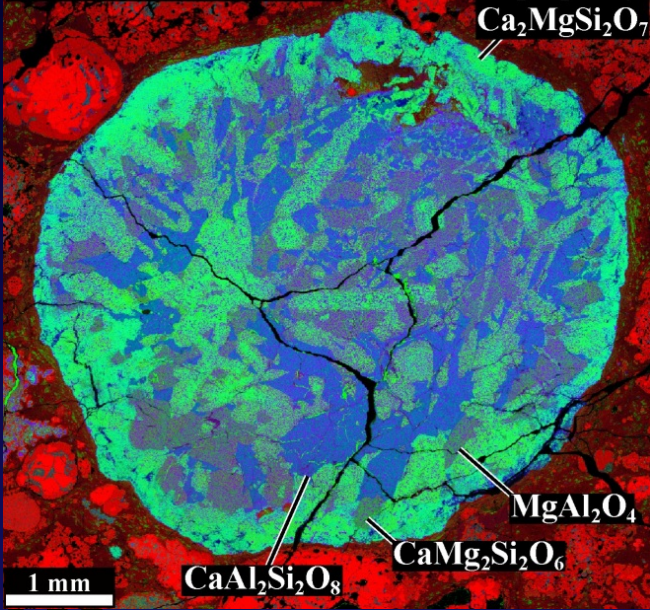
# II. Ca,Al-rich Inclusions (CAIs)

**fine-grained CAI**

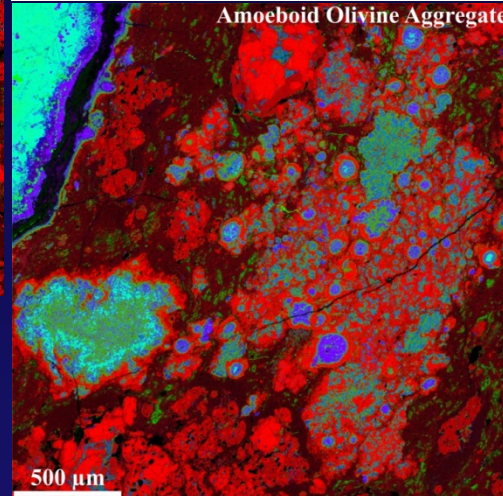
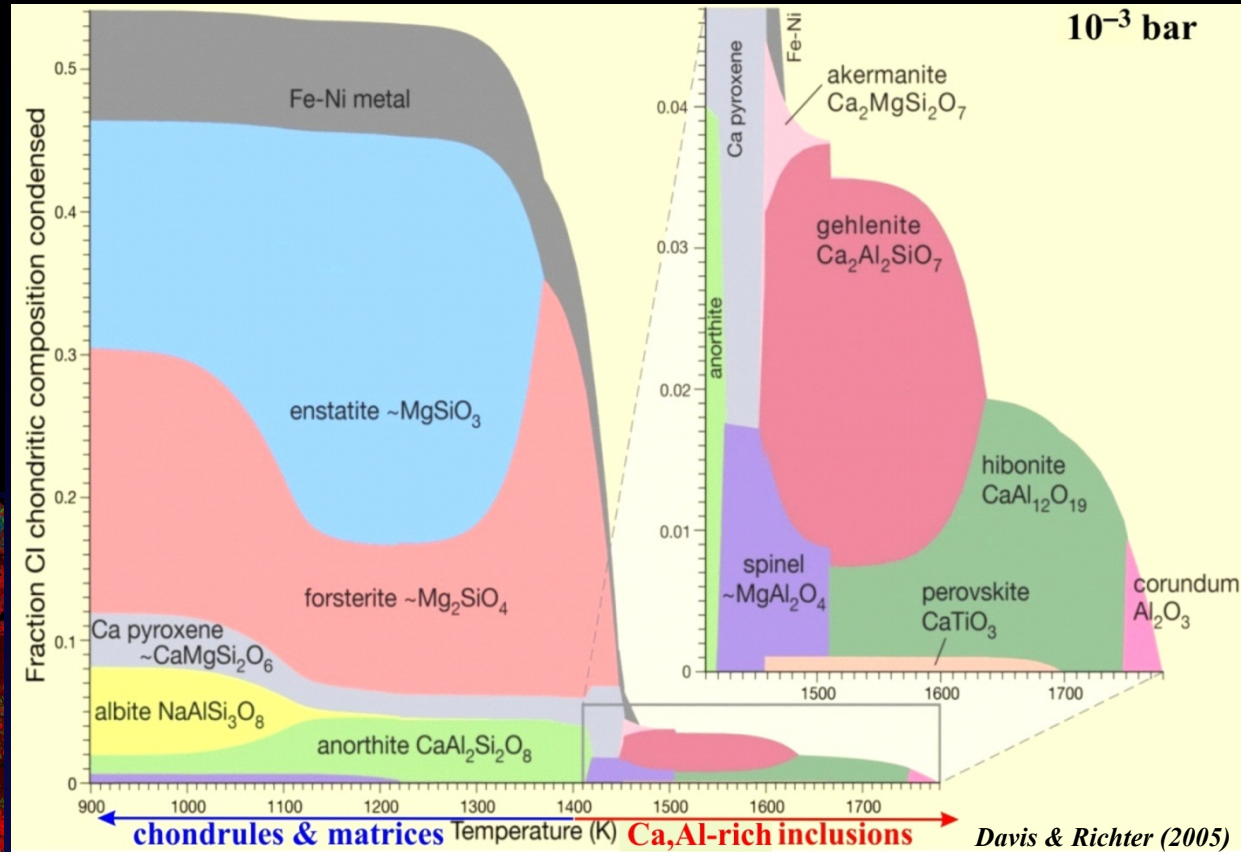


**1 mm**

**remelted condensates**



**1 mm**

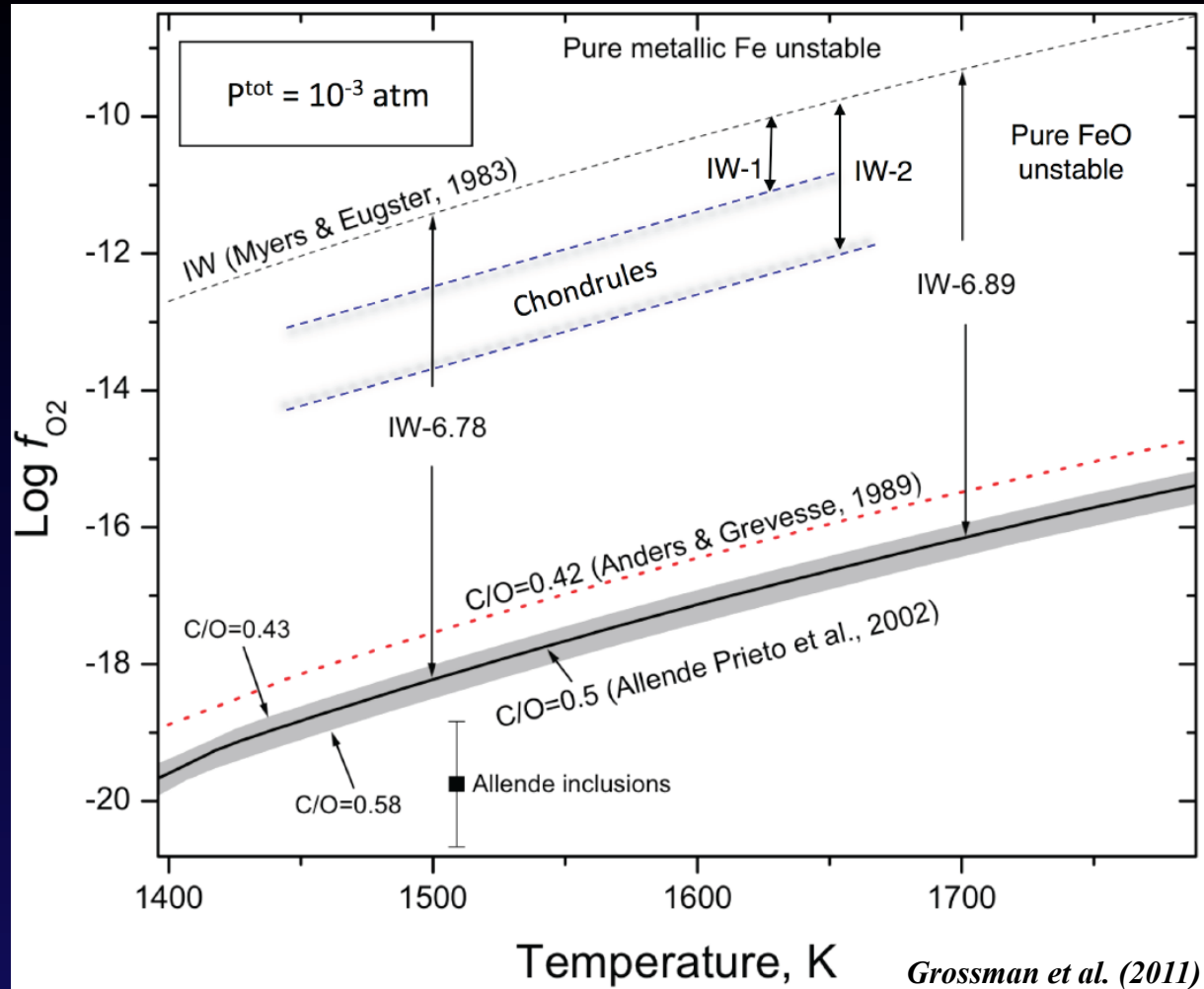


**500 μm**

- by multistage evaporation, condensation, aggregation, & melting of solids in region(s) with ambient  $T > 1400$  K
- oldest Solar System solids
- CV CAIs:  $4567.3 \pm 0.16$  Myr  $\equiv$  age of Solar System or time 0

• cm-sized CAIs formed at time 0!

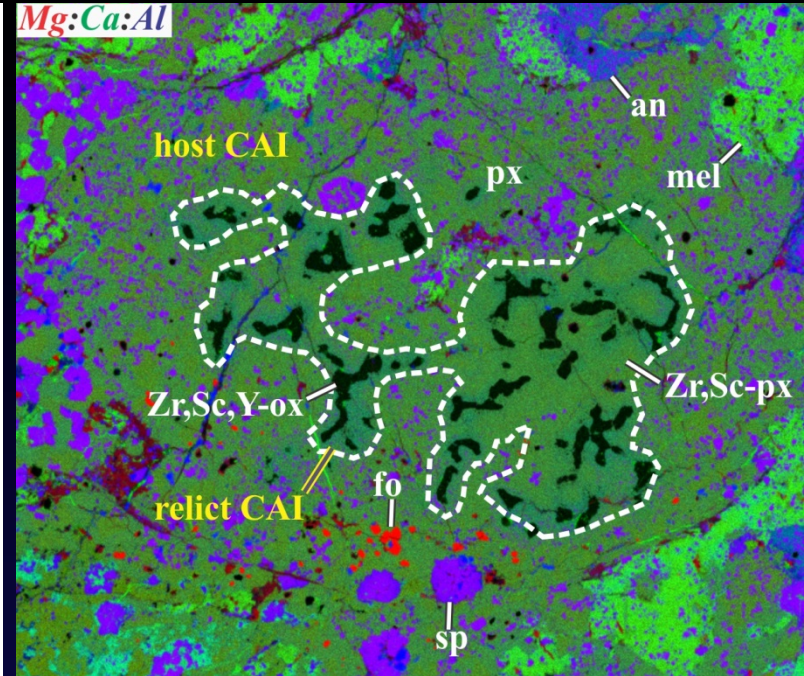
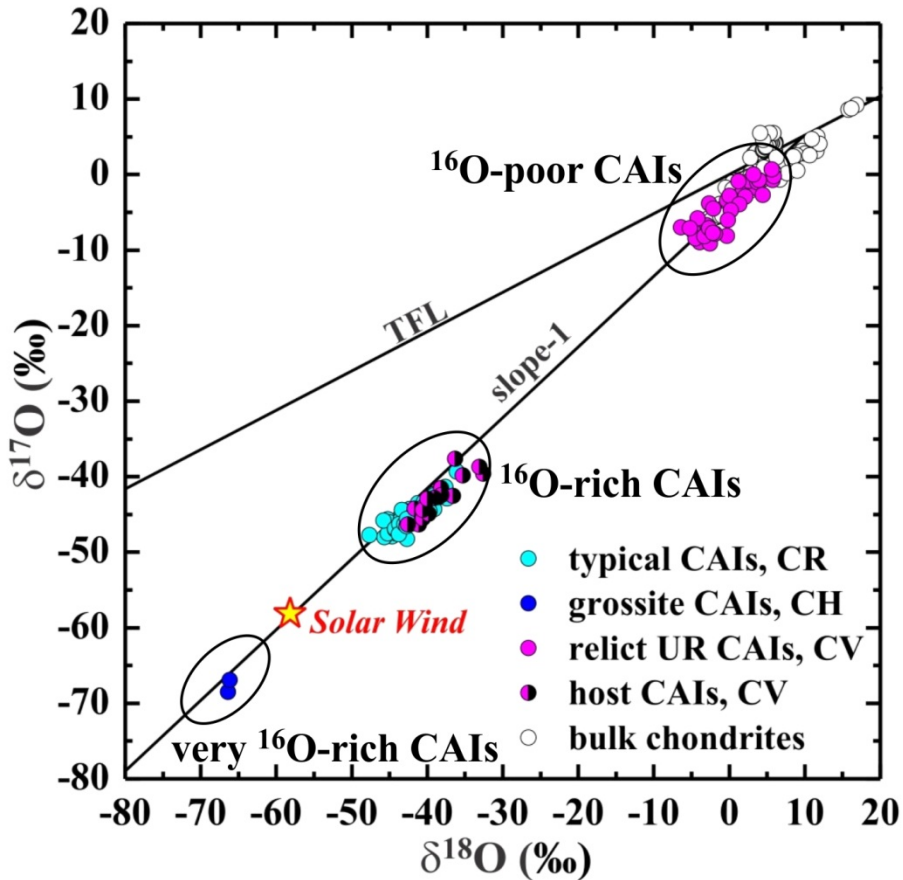
# Redox conditions in CAI-forming region(s)



- CAI minerals contain no oxidized Fe ( $\text{Fe}^{2+}$  or  $\text{Fe}^{3+}$ ) & have high  $\text{Ti}^{3+}/\text{Ti}^{4+}$  ratio  
→ formed in highly-reducing, gas-dominated region(s) with a solar  $\text{H}_2\text{O}/\text{H}_2 \sim 5 \times 10^{-4}$



# Oxygen-isotope compositions of CAIs



- relict  $^{16}\text{O}$ -depleted ultrarefractory CAI inside  $^{16}\text{O}$ -rich host CAI

- most CAIs:  $^{16}\text{O}$ -rich, close to solar wind value
- some CAIs:  $^{16}\text{O}$ -enriched or  $^{16}\text{O}$ -depleted relative solar wind
- early generation of isotopically distinct oxygen reservoirs in PPD
- mechanism of generation of these reservoirs is not understood (CO self-shielding?)
- O-isotope compositions of primordial dust & gas are not known

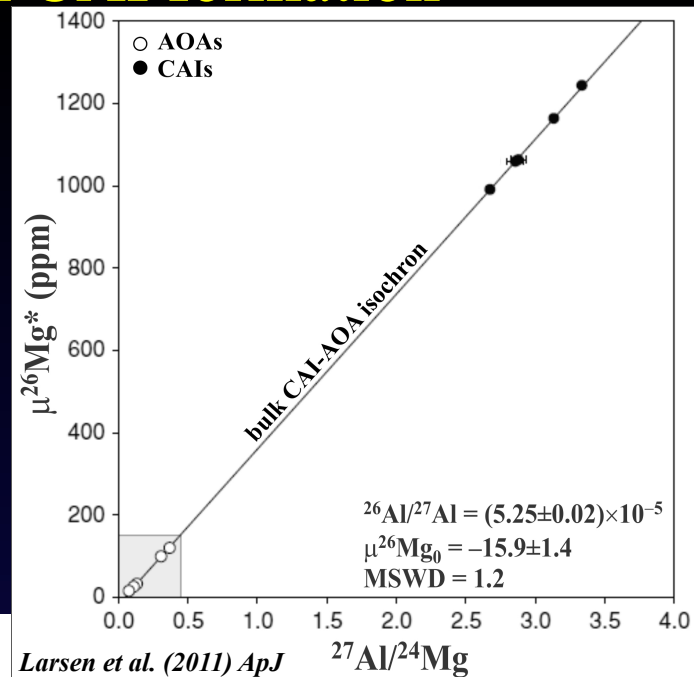
# Short-lived radionuclides ( $^{53}\text{Mn}$ , $^{60}\text{Fe}$ , $^{182}\text{Hf}$ , $^{10}\text{Be}$ , $^{26}\text{Al}$ , $^{41}\text{Ca}$ )

- $^{53}\text{Mn}$  ( $t_{1/2} \sim 3.7$  Myr),  $^{60}\text{Fe}$  (2.6 Myr),  $^{182}\text{Hf}$  (9 Myr): uniformly distributed in PPD & inherited from MC
    - $^{53}\text{Mn}$  &  $^{182}\text{Hf}$  are used for chronology of SS processes
  - $^{10}\text{Be}$  ( $t_{1/2} \sim 1.5$  Myr),  $^{26}\text{Al}$  (0.7 Myr), &  $^{41}\text{Ca}$  (0.1 Myr): heterogeneous among CAIs
    - $^{10}\text{Be}$ : energetic particle irradiation near protoSun
      - formation of CAIs near protoSun
    - $^{26}\text{Al}$  &  $^{41}\text{Ca}$  correlate with each other, but do not correlate with  $^{10}\text{Be}$ ;
    - $^{26}\text{Al}$  in PPD is too high to be explained by irradiation *Dupra & Tatischeff (2007) ApJ*
      - external, stellar origin
      - Solar System formed near massive star(s) (SN\*, AGB, Wolf-Rayet)
- \*previously inferred high abundance of  $^{60}\text{Fe}$  requiring SN source has not been confirmed (*check poster #43 by M. Telus*)

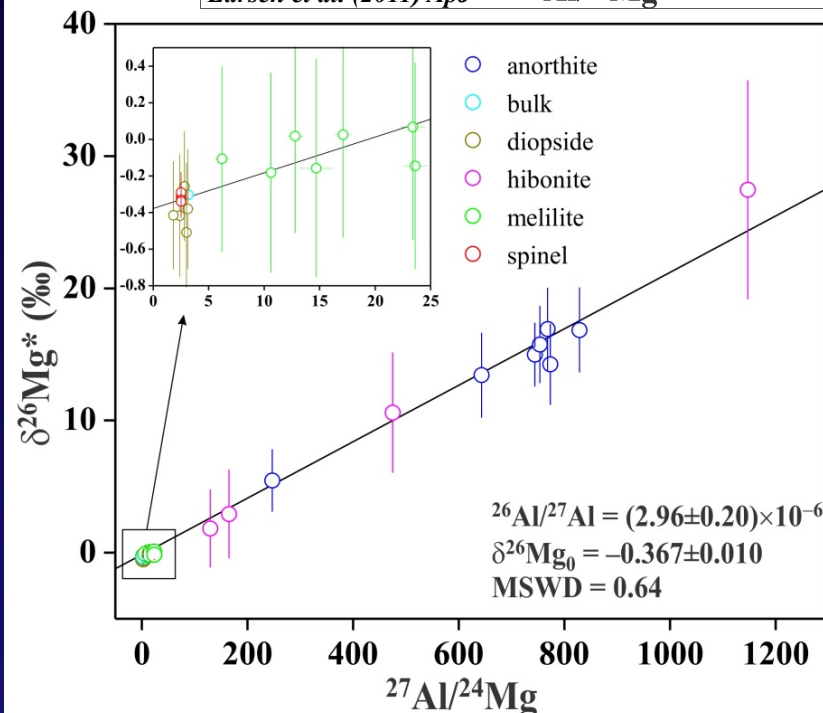
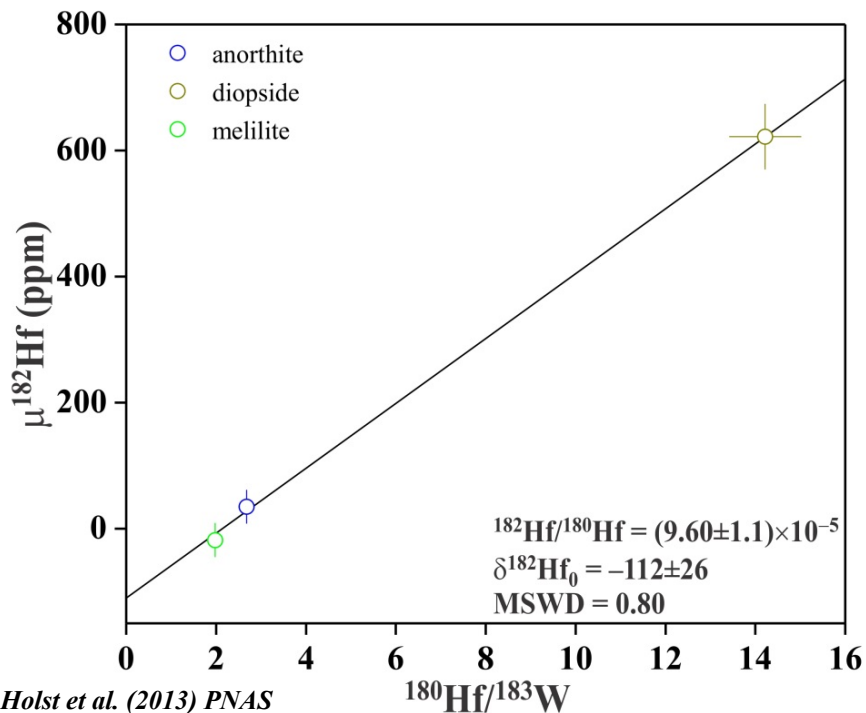


# $^{26}\text{Al}$ heterogeneity during epoch of CAI formation

- CV CAIs:  $^{26}\text{Al}/^{27}\text{Al} = (5.25 \pm 0.02) \times 10^{-5}$ , “canonical”
  - uniform distribution of  $^{26}\text{Al}$  in PPD
  - brief ( $< 0.002$  Myr) duration of CAI formation
- incorrect:
  - some CAIs:  $^{26}\text{Al}/^{27}\text{Al} < 5 \times 10^{-6}$  & formed before or contemporaneously with canonical CAIs
    - $\rightarrow$   $^{26}\text{Al}$  heterogeneity in protosolar MC
    - $\rightarrow$  recent injection of  $^{26}\text{Al}$  into MC core or PPD



## CAIs with diff. $^{26}\text{Al}/^{27}\text{Al}$ have similar $^{182}\text{Hf}/^{180}\text{Hf}$

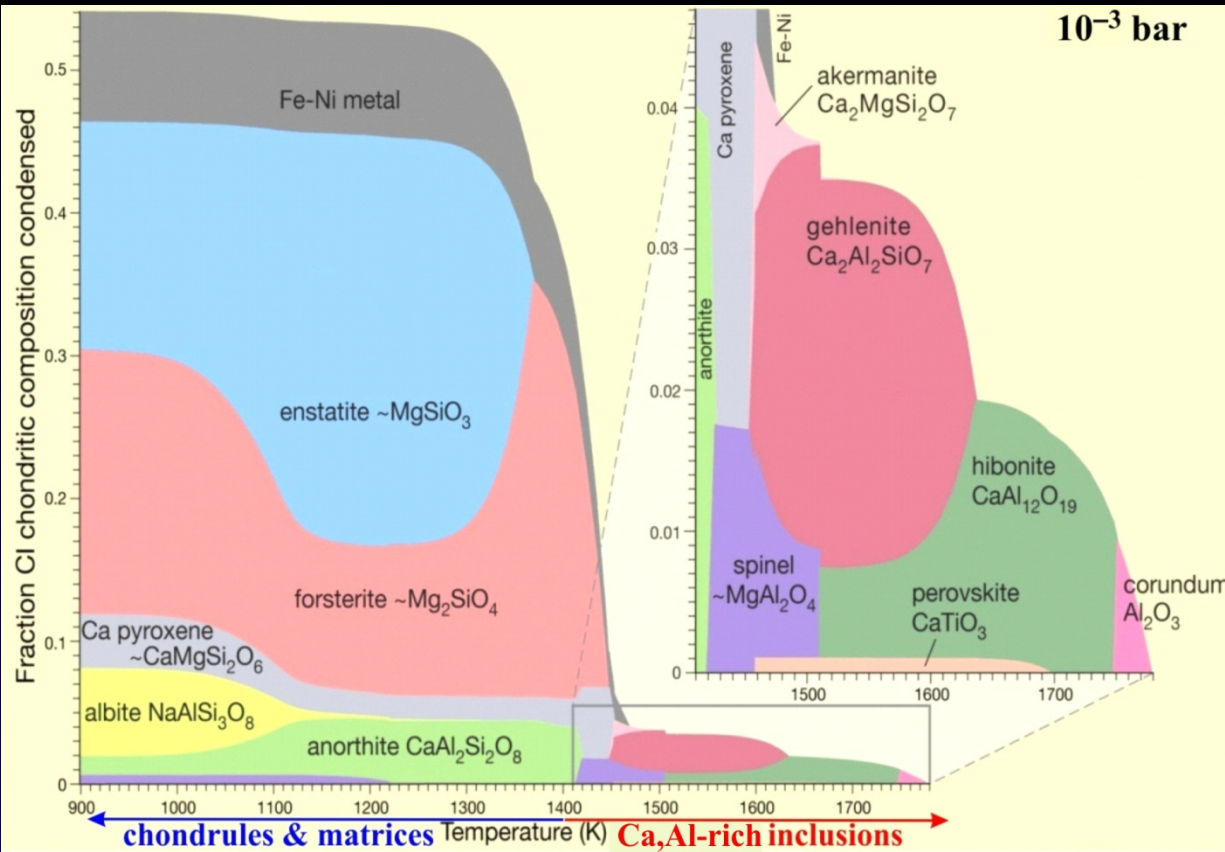


# CAIs: Summary

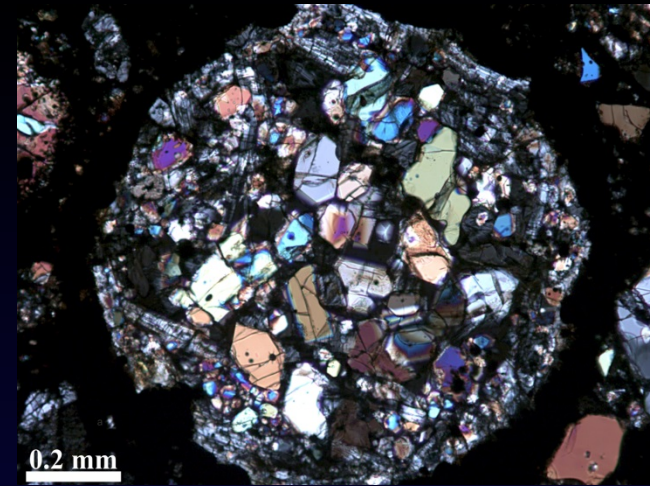
- earliest SS solids dated:  $\sim 4567.3$  Ma  $\equiv$  age of SS  $\equiv$  time 0 in cosmochemistry
- evaporation, condensation, aggregation, irradiation ( $^{10}\text{Be}$ ) &  $\pm$ melting processes in a gas of solar composition (reduced &  $^{16}\text{O}$ -rich) in region(s) with ambient  $T > 1400\text{K}$
- early generation of isotopically distinct oxygen reservoirs in PPD
- heterogeneous distribution of  $^{26}\text{Al}$  &  $^{41}\text{Ca}$  in PPD
  - recent injection of  $^{26}\text{Al}$  into  $^{26}\text{Al}$ -poor MC core by massive star(s)
  - duration of CAI formation is not known & cannot be inferred from  $^{26}\text{Al}$
  - distribution of  $^{26}\text{Al}$  in PPD cannot be inferred from CAIs
- present in all chondrite groups & in a comet *81P/Wild 2*
  - after formation were removed from hot region & dispersed throughout PPD



# III. Chondrules & fine-grained matrices



*Davis & Richter (2005)*



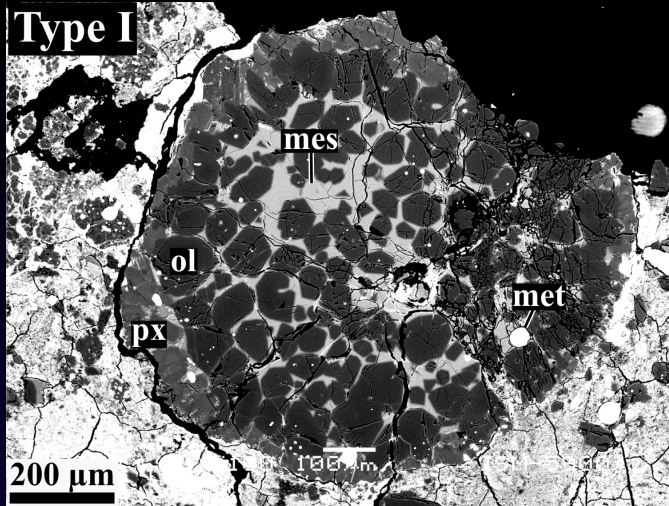
*courtesy of D. Lauretta*

mm-sized molten & rapidly solidified objects once freely floating in space

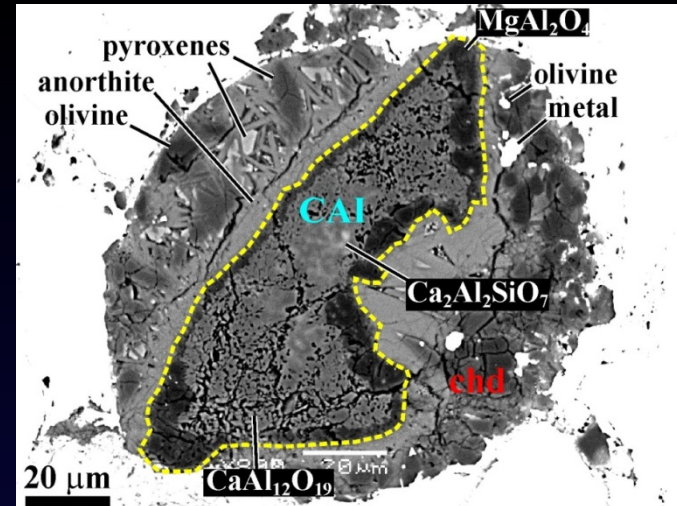
- less refractory than CAIs
- contain ferromagnesian silicates (Fe<sup>2+</sup>) & abundant volatiles (Na, K, S)



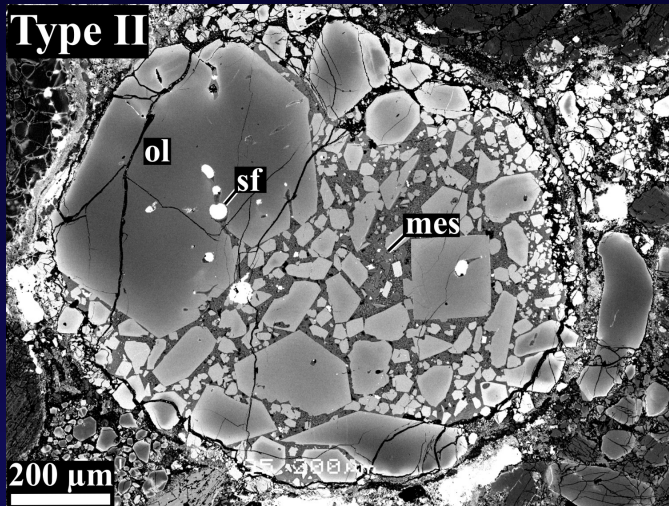
# Porphyritic textures & relict grains: Incomplete melting



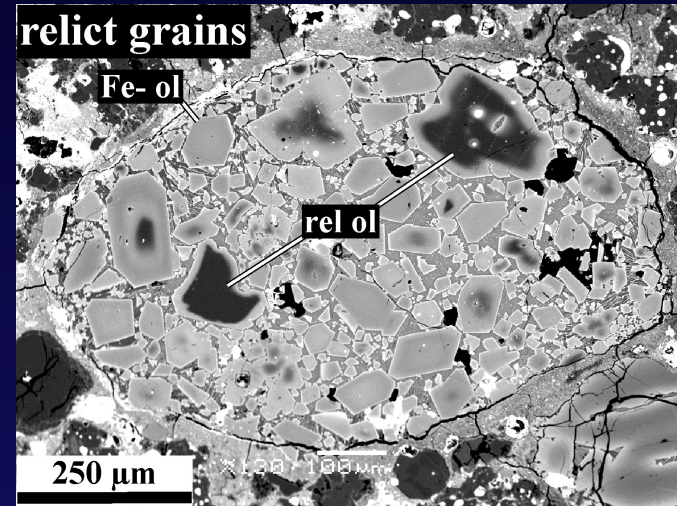
magnesian porphyritic



relict CAI



ferroan porphyritic

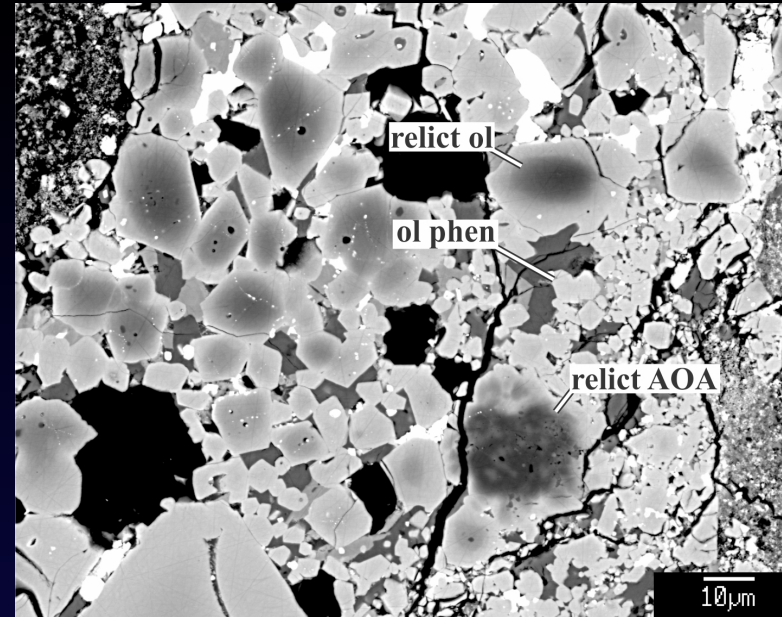
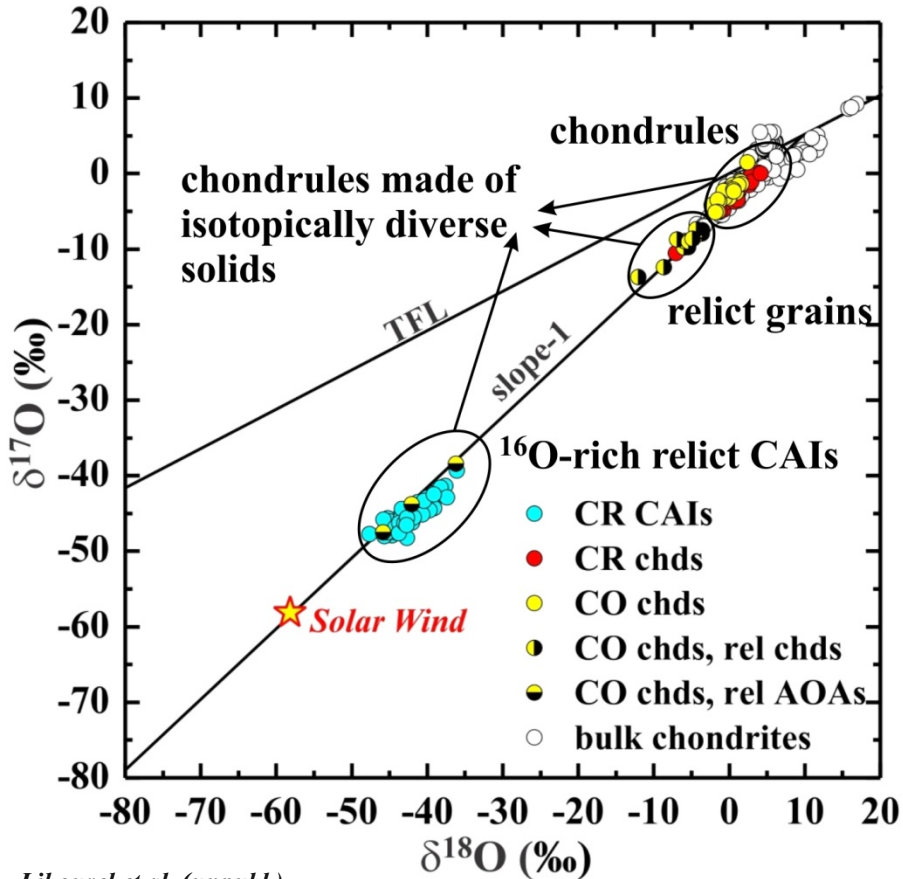


relict chondrule fragments

- melting (often incomplete) of solid precursors, including fragments of earlier formed chondrules & CAIs



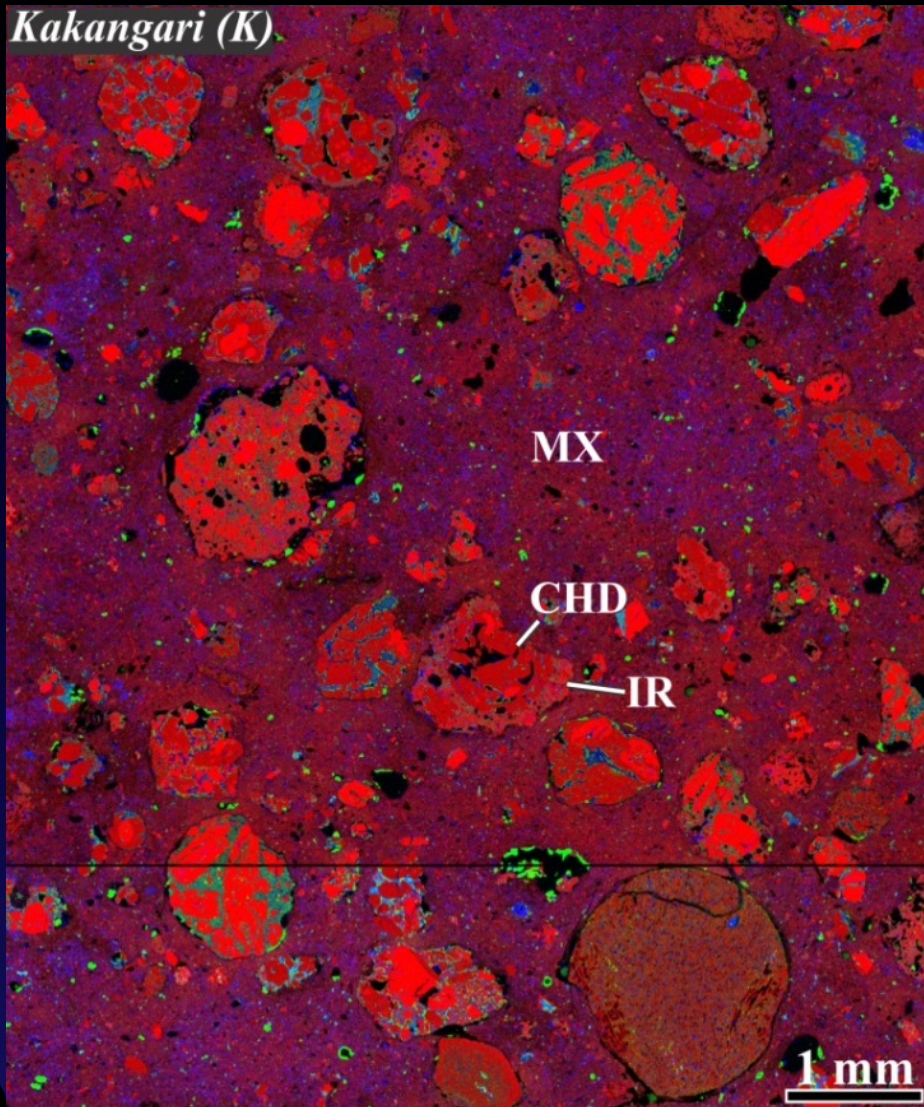
# Oxygen-isotope compositions of chondrules



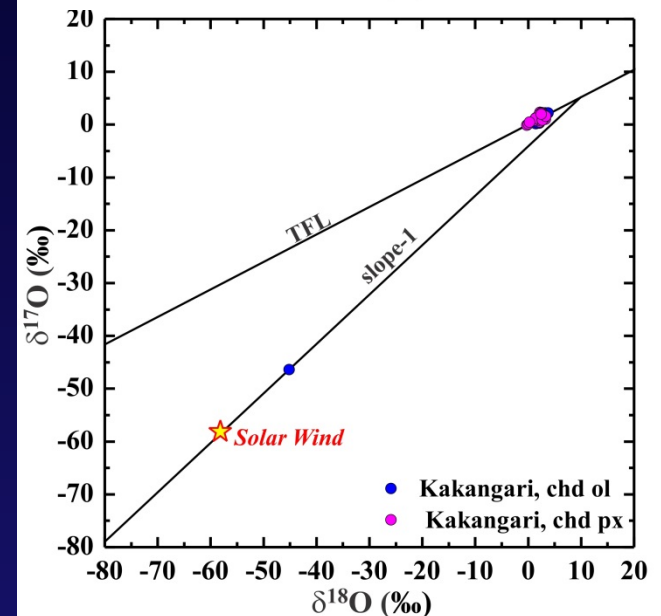
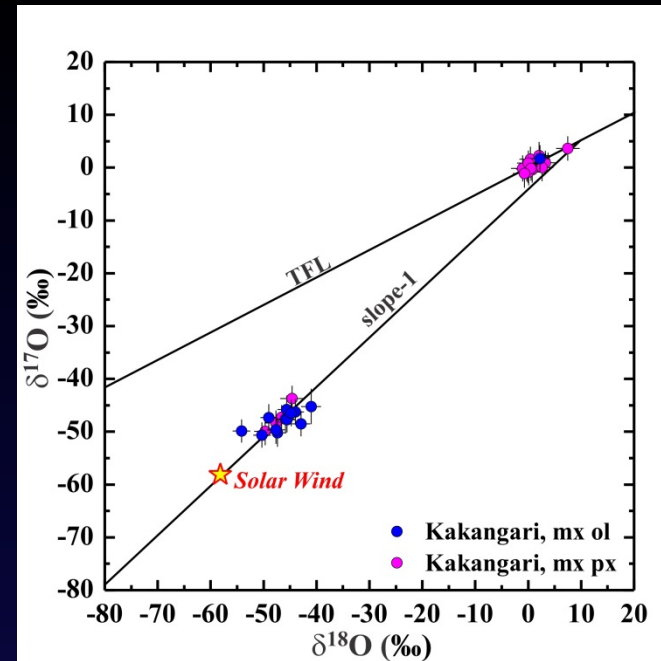
relict AOA & relict chondrule fragments in a ferroan porphyritic chondrule

- chondrules are  $^{16}\text{O}$ -depleted relative to CAIs
  - relict grains  $^{16}\text{O}$ -enriched relative host chondrules
- chondrules formed by melting of isotopically diverse precursors in  $^{16}\text{O}$ -depleted gaseous reservoir

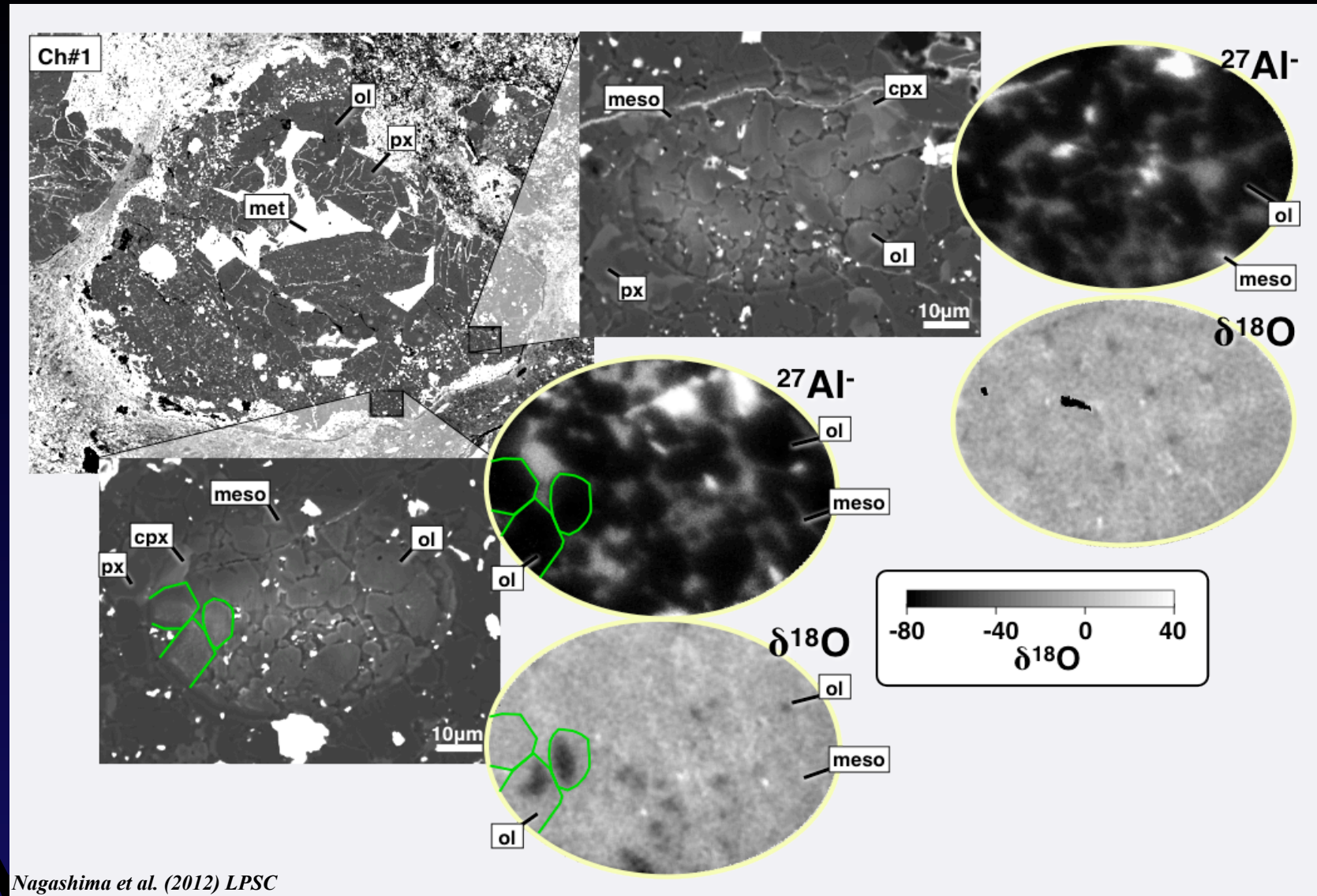
# Chondrule-matrix relationship: Insights from oxygen isotopes



- matrix & chondrules in Kakangari contain isotopically similar olivine & pyroxene



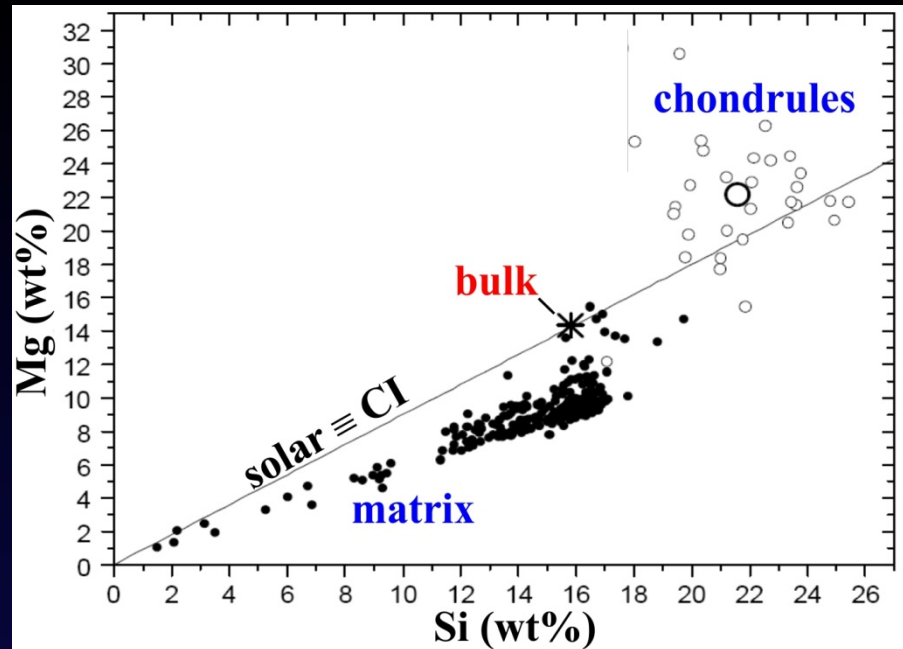
# Chondrule-matrix relationship: Insights from oxygen isotopes



*Nagashima et al. (2012) LPSC*

- matrix & igneous rims around Kakangari chondrules contain abundant  $^{16}\text{O}$ -rich grains  
→ matrix grains were among chondrule precursors

# Chondrule-matrix relationship: Insights from bulk chemistry



*Klerner & Palme (1999) MAPS*

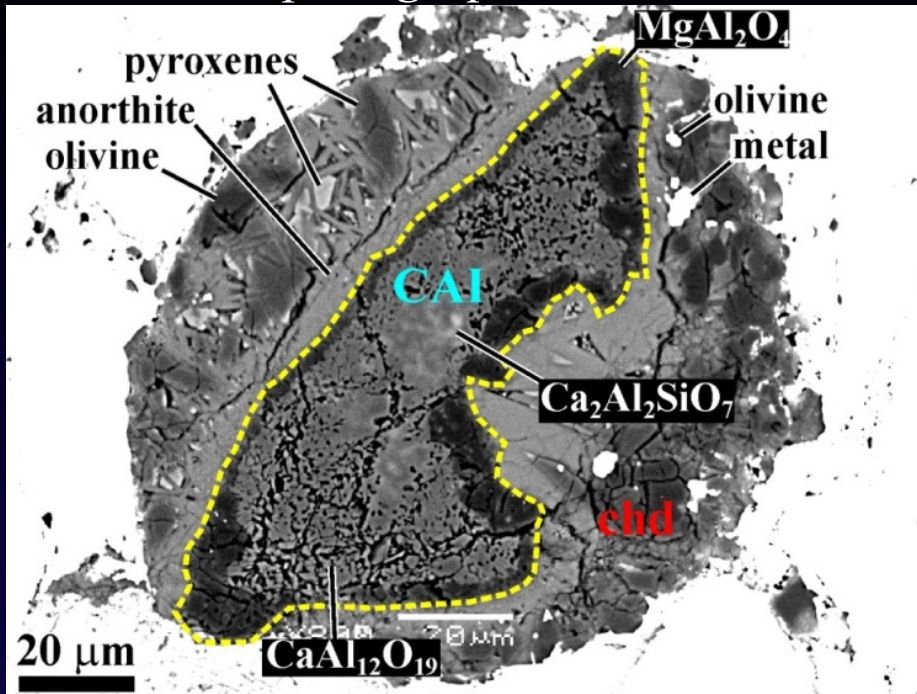
- matrix & chondrules in a chondrite group are chemically complementary: e.g., Mg/Si ratio in Renazzo, CR chondrite

matrix	$0.65 \pm 0.11$
chondrules	$1.03 \pm 0.20$
bulk chondrite	0.91
solar Mg/Si	0.90

→ matrix & chondrules formed in the same nebular regions throughout the PPD, contrary to X-wind model suggesting chondrules formed near the protoSun & were transported to 1-4 AU where they accreted together with thermally unprocessed matrices (*Shu et al. 1996, 1997, 2001*)

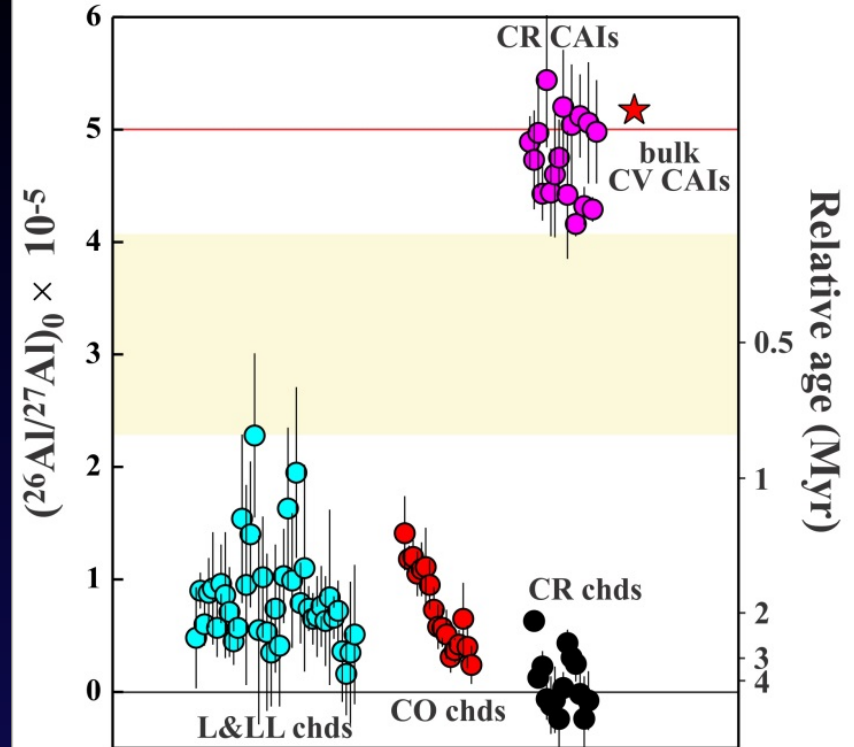
# Chondrules: Relative & absolute chronology

based on petrographic observations



- relict CAIs in chondrules
- CAIs predated chondrules
- age difference cannot be inferred

based on short-lived radionuclide  $^{26}\text{Al}$



assumption: uniform  $^{26}\text{Al}/^{27}\text{Al}$  in PPD

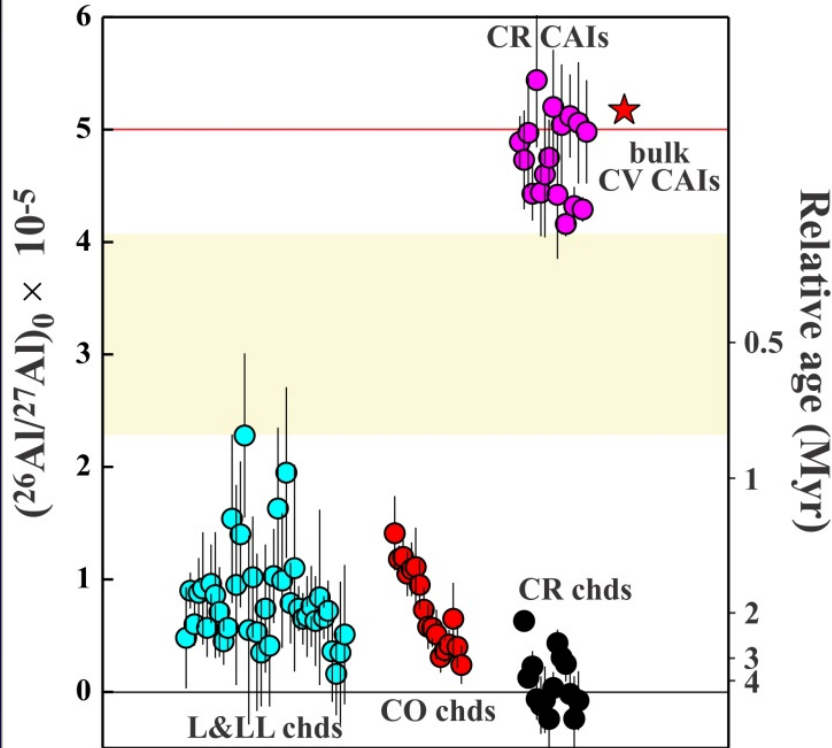
Krot et al. (2009) GCA

- if  $(^{26}\text{Al}/^{27}\text{Al})_0$  in PPD was uniform,  $\sim 5 \times 10^{-5}$
- 1 Myr age gap between CAIs & chondrules
- chondrule formation lasted for  $\sim 3$  Myr
- life-time of PPD is at least 4 Myr



# Chondrules: Relative $^{26}\text{Al}$ - $^{26}\text{Mg}$ & absolute U-Pb chronology

assumption: uniform  $^{26}\text{Al}/^{27}\text{Al}$  in PPD

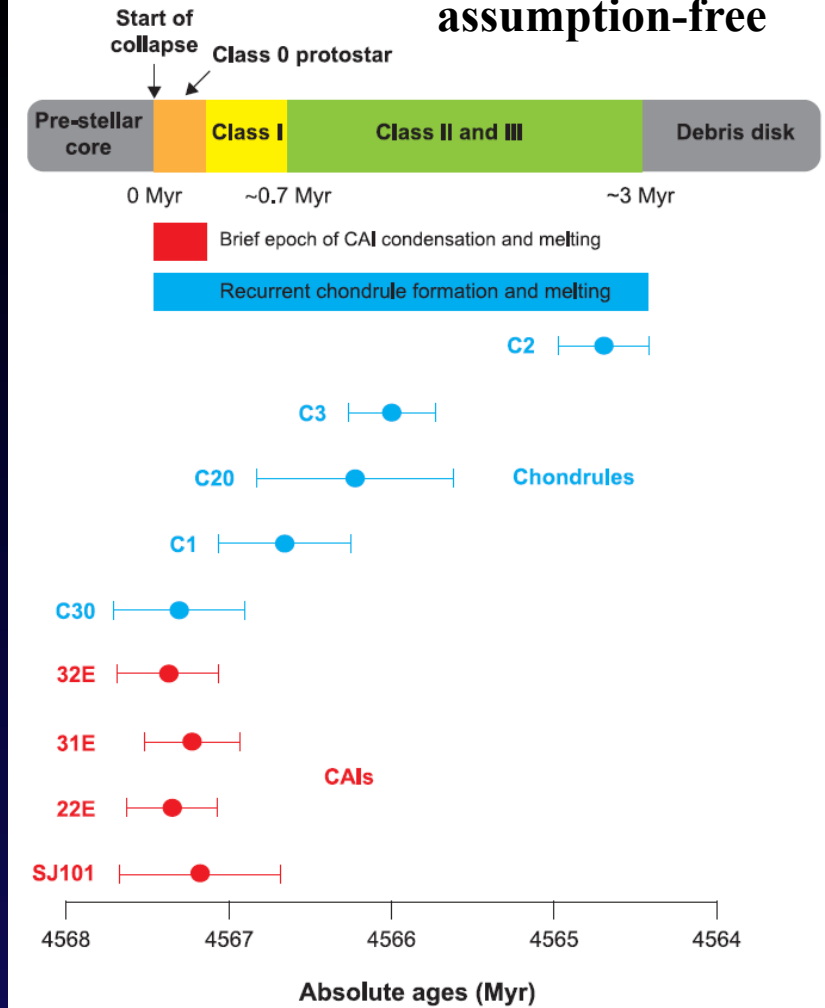


Krot et al. (2009) GCA

## Relative $^{26}\text{Al}$ - $^{26}\text{Mg}$ chronology:

- 1 Myr age gap between CAIs & chondrules
- chondrule formation lasted for  $\sim 3$  Myr
- at least 4 Myr PPD life-time

assumption-free

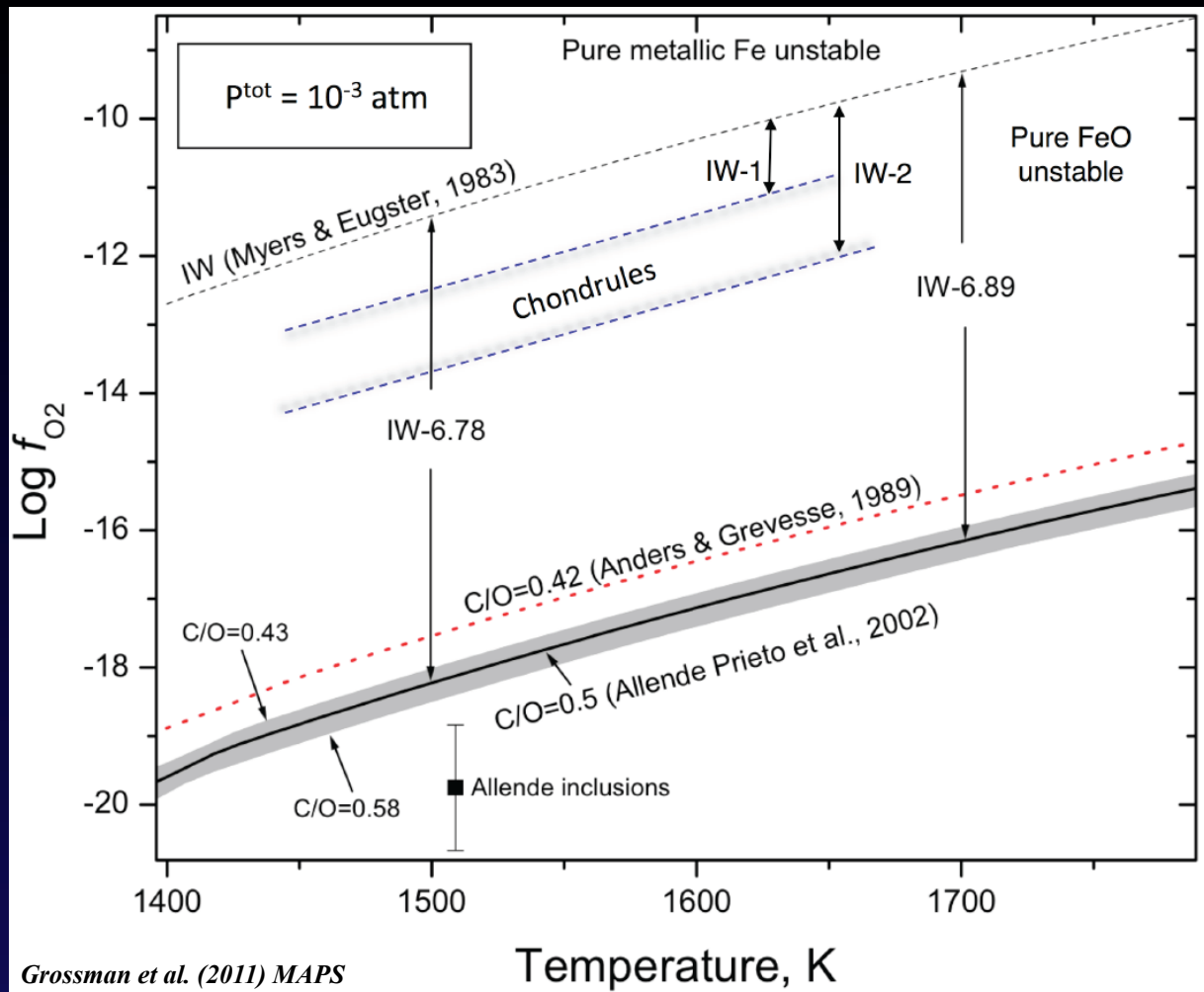


Connelly et al. (2012) Science

## Absolute U-Pb chronology:

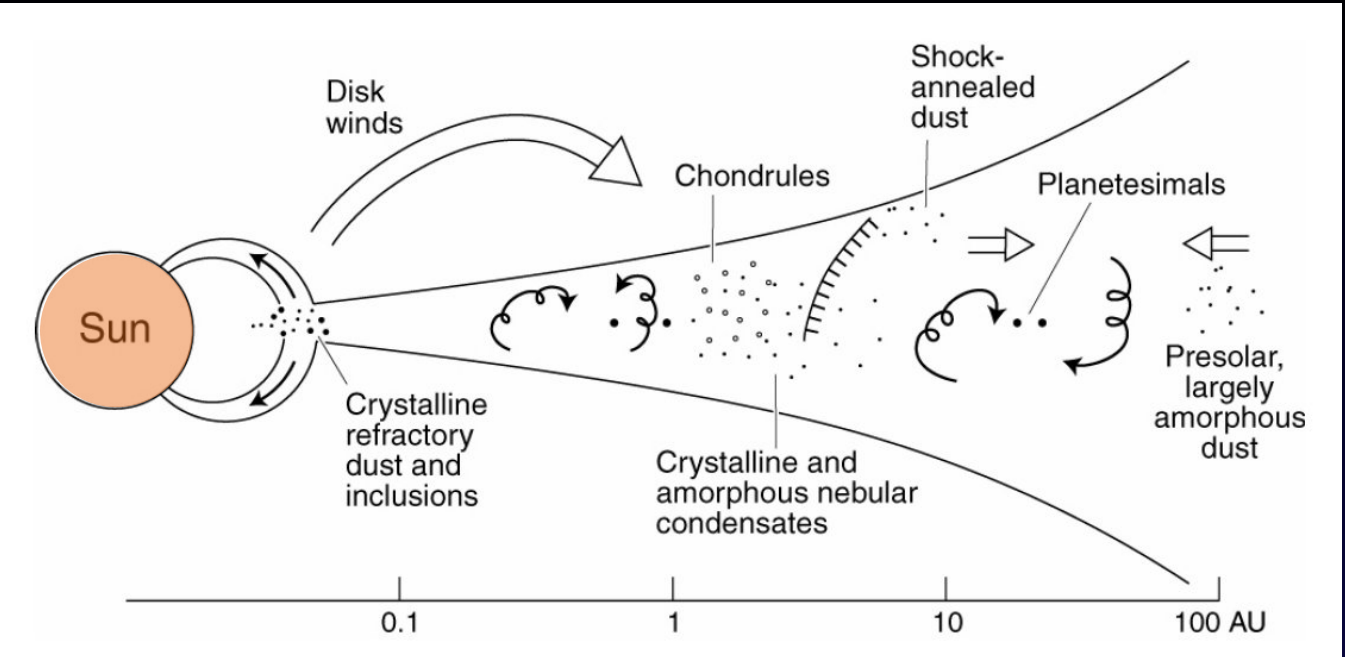
- chondrule formation started contemporaneously with CAIs & lasted for at least 3 Myr

# Chondrules formed under oxidizing conditions

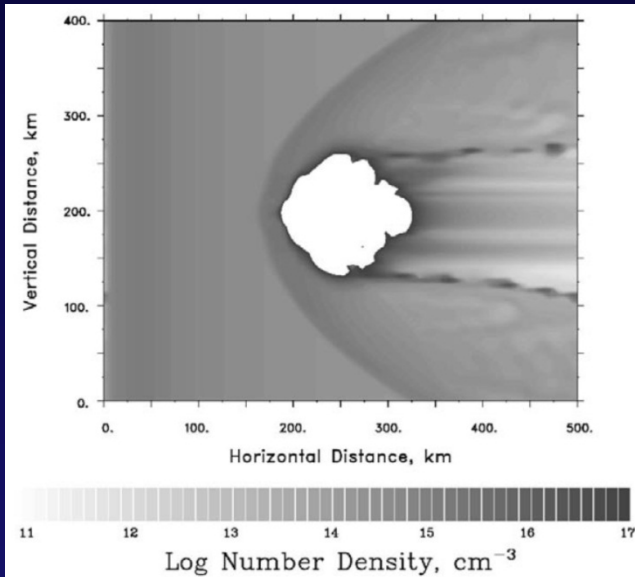


- under much more oxidizing (up to IW-1) conditions than CAIs (up to IW-7)
- high partial pressures of Si, Na, Mg, Fe, & S (were not lost from chondrule melts)
- formed under non-solar conditions ( $\text{D/G} > 10^4 \times \text{solar}$ ;  $\text{H}_2\text{O}/\text{H}_2 > 10^2 \times \text{solar}$ )

# Models of chondrule formation: Shock waves, impacts, lightning ...

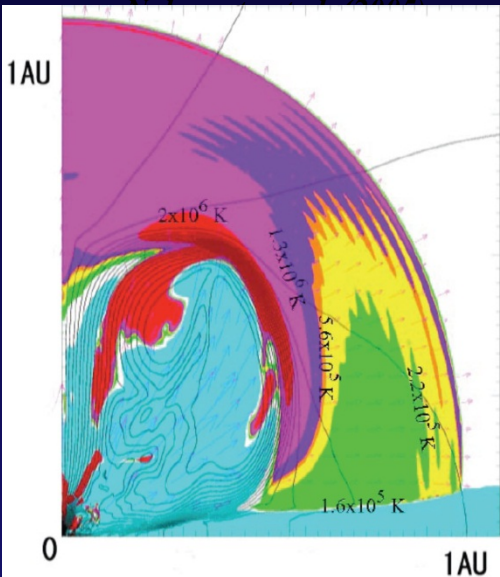


by fast moving planetesimals



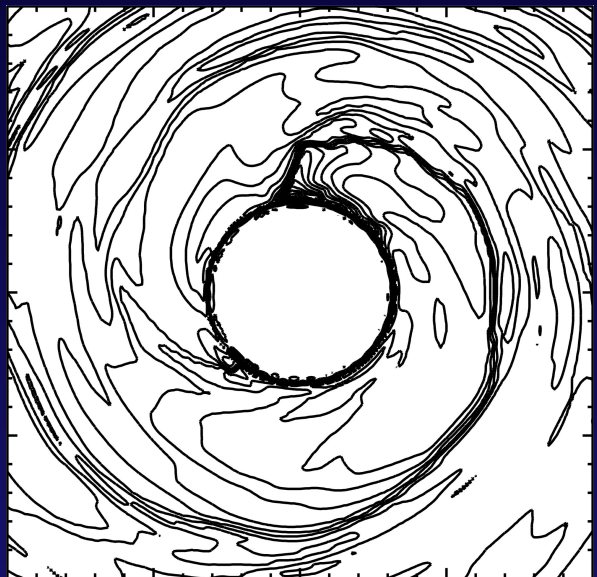
Hood et al. (2005)

by X-ray flares



Nakamoto et al. (2005)

by gravitational instability



Boss & Durisen (2005)



# Chondrules & fine-grained matrices: Summary

- in isotopically distinct regions of inner disk dominated by  $^{16}\text{O}$ -poor dust & gas
  - at lower ambient T (<650 K) & highly non-solar dust/gas &  $\text{H}_2\text{O}/\text{H}_2$  ratios
  - rapid heating (up to  $1600^\circ\text{C}$ ) & cooling ( $1\text{-}1000^\circ\text{C hr}^{-1}$ ) of isotopically & mineralogically diverse solid precursors (fragments of earlier formed chondrules & CAIs, & matrix)
  - most matrix was thermally processed during chondrule & CAI formation
  - formation mechanisms are not understood: shock waves, impacts, lightning, ...
  - started contemporaneously with CAIs & lasted  $\sim 3\text{-}4$  Myr
- life-time of PPD  $\sim 3\text{-}4$  Myr
- duration of accretion of chondrite asteroids  $\sim 3\text{-}4$  Myr
- chondrule formation may have been rapidly followed by chondrite accretion



# IV. Aqueous activity on chondrite parent bodies

*Carbonaceous*

*Enstatite*

*Ordinary*

*Other*

*Ungrouped*

CI CM CR CV CK CO

CH CB

EH EL

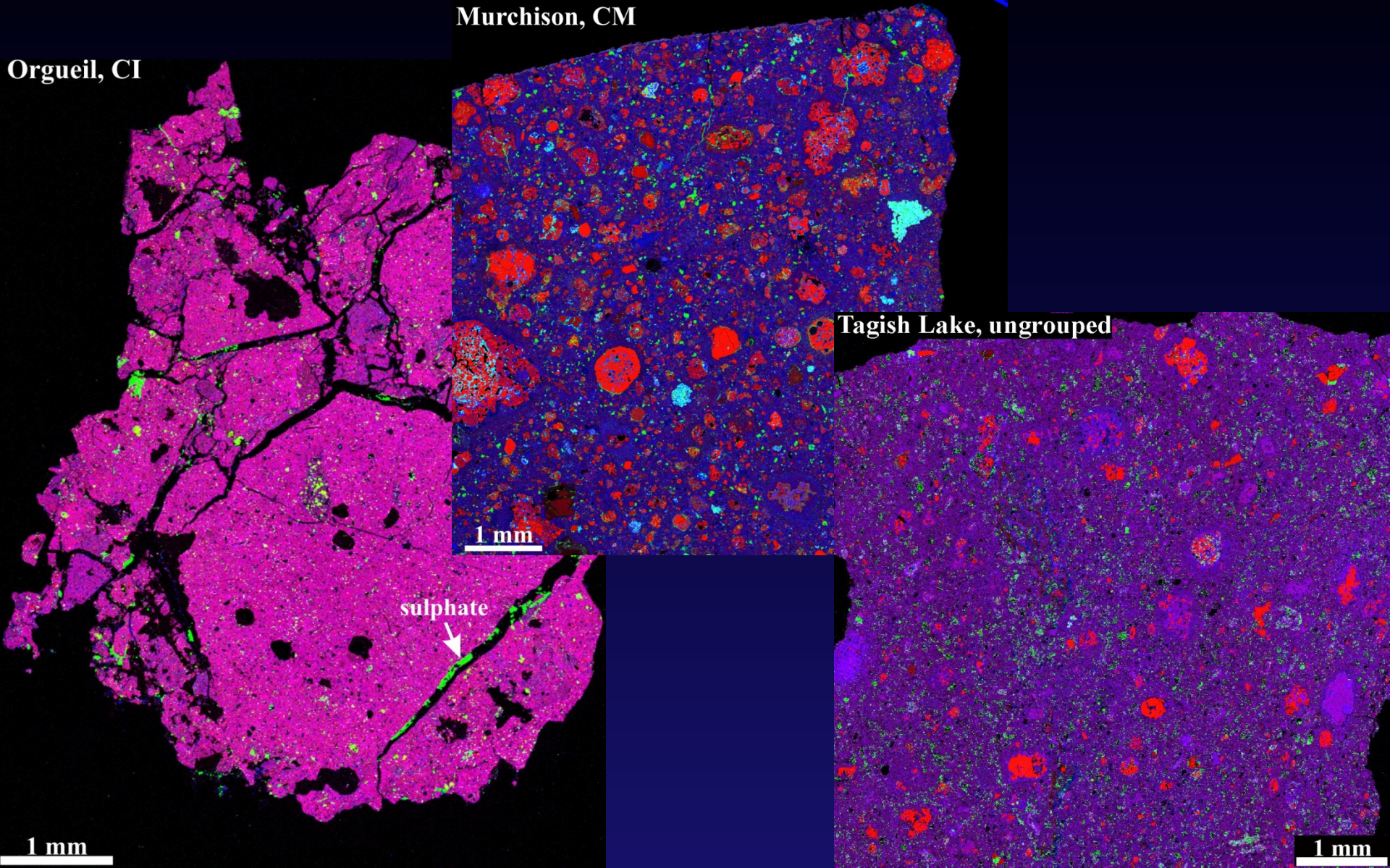
H L LL

K R

Tagish Lake

Murchison, CM

Orgueil, CI



1 mm

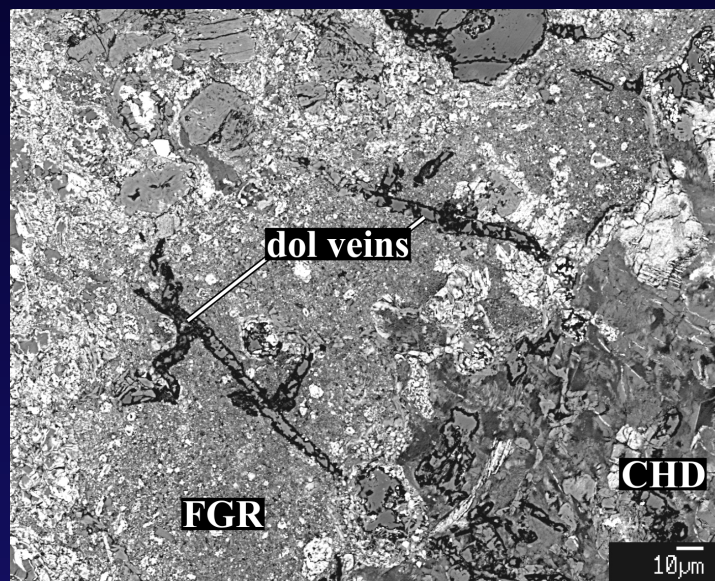
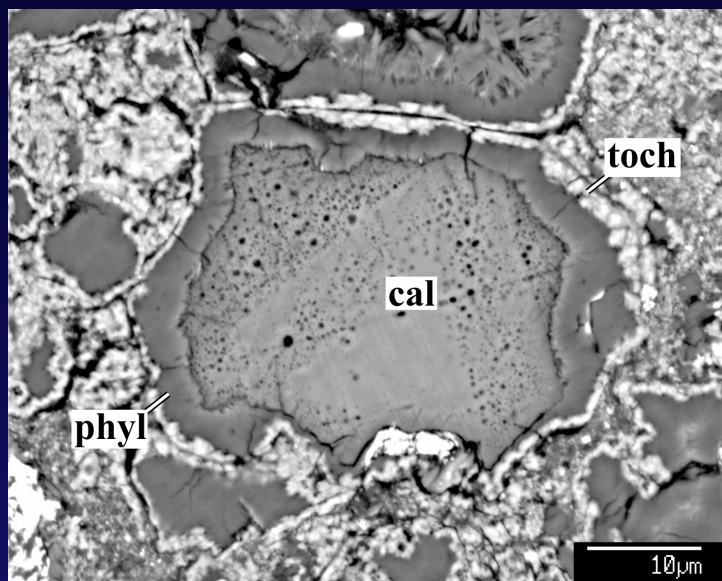
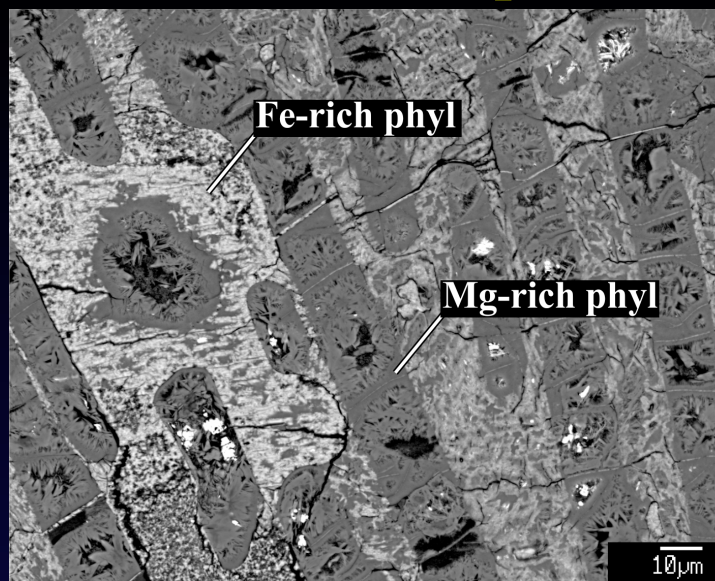
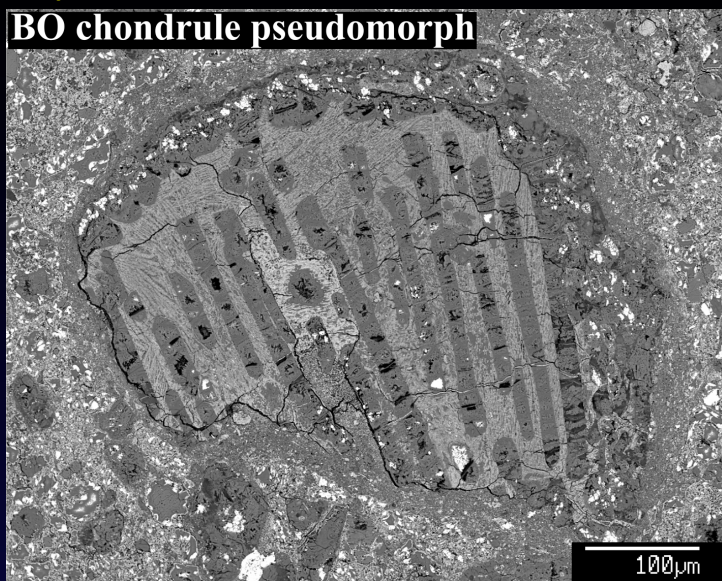
sulphate

Tagish Lake, ungrouped

1 mm

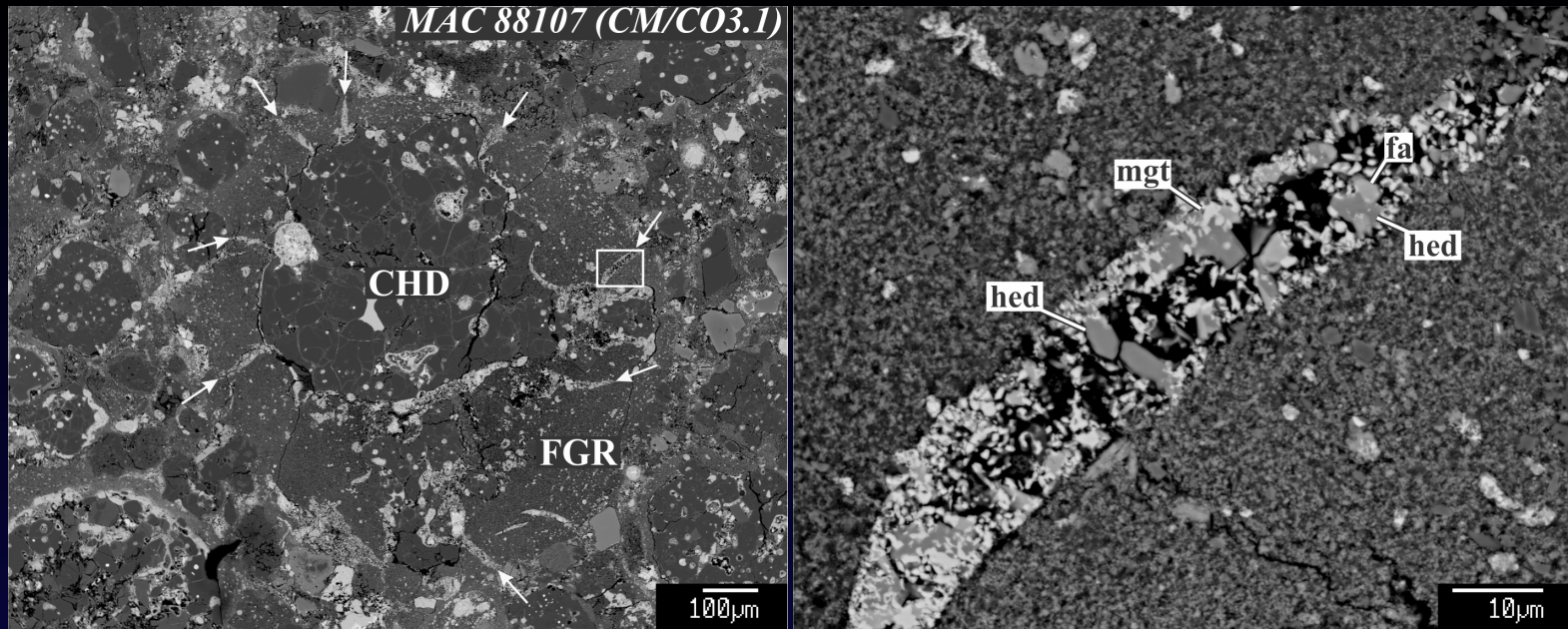
1 mm

# Phyllosilicates, carbonates, veins, chondrule pseudomorphs



- aqueous alteration occurred on chondrite asteroids, not in the nebula
- CMs, CIs, & CRs: low-T aq. alteration ~25-100°C at high W/R vol. ratio: 0.4-1

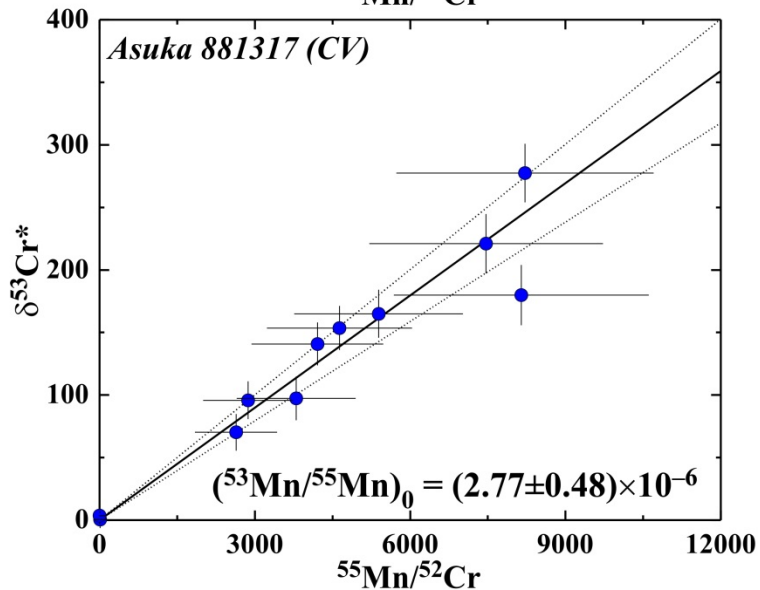
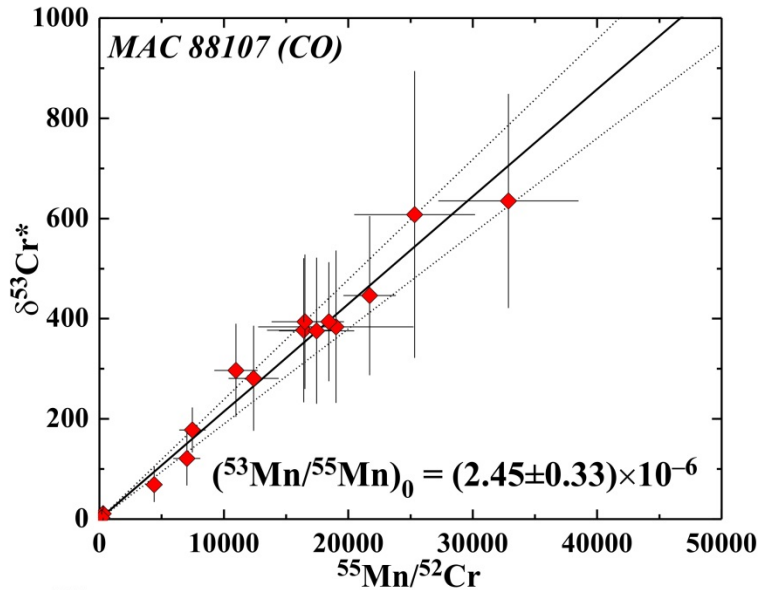
# Fayalite-hedenbergite-magnetite veins



- CV, CO, CK, H, L, LL, & R chondrites:  $T \sim 100\text{-}200^{\circ}\text{C}$  & low  $W/R < 0.2$
- aqueous alteration under highly-oxidizing conditions
- phyllosilicates, fayalite ( $\text{Fe}^{2+}_2\text{SiO}_4$ ), magnetite ( $\text{Fe}^{2+}\text{Fe}_2^{3+}\text{O}_4$ ), hedenbergite  $\text{CaFe}^{2+}\text{Si}_2\text{O}_6$ , andradite  $\text{Ca}_3\text{Fe}_2^{3+}\text{Si}_3\text{O}_{12}$



# Chronology of aqueous alteration



$^{53}\text{Mn}$ - $^{53}\text{Cr}$  system:  $^{53}\text{Mn} \rightarrow ^{53}\text{Cr}$ ,  $t_{1/2} \sim 3.7$  Myr

$^{53}\text{Mn}$  uniformly distributed in PPD with the  $(^{53}\text{Mn}/^{55}\text{Mn})_0 = 6 \times 10^{-6}$  (Kleine et al., 2012, GCA)

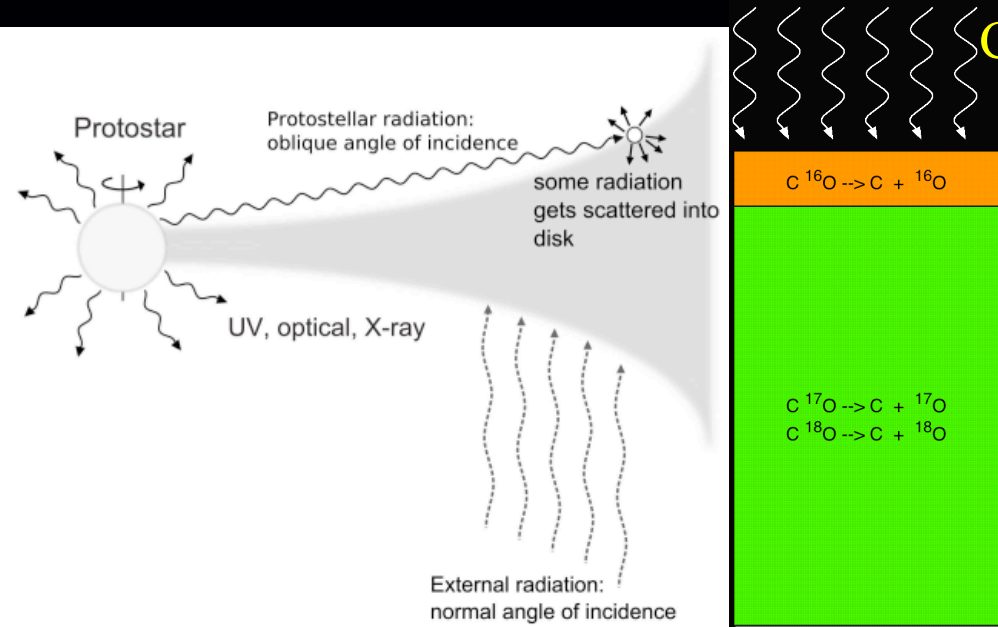
Doyle et al. (2013) LPSC

- fayalite in CVs : 3.7 Myr after  $t_0$
- fayalite in COs : 4.4 Myr after  $t_0$

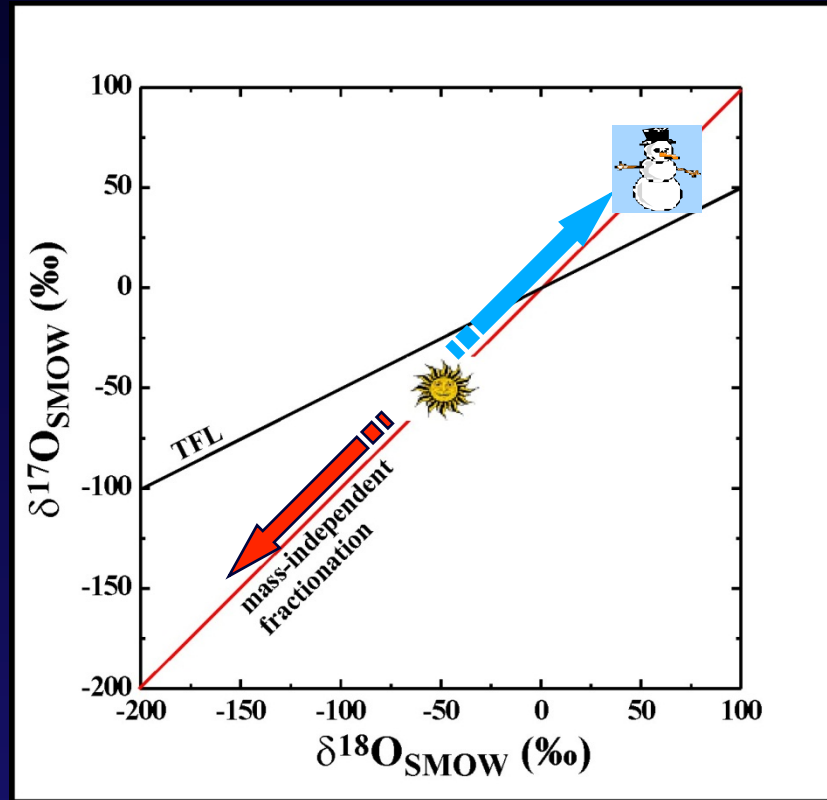
Fujiya et al. (2012, 2013) EPSL, Nature Comm.

- carbonates in CMs :  $\sim 4$  Myr after  $t_0$
  - carbonates in CIs :  $\sim 3.5$  Myr after  $t_0$
- aqueous alteration on CC parent asteroids started shortly after accretion
- chondrites formed near the snow line
- position of snow line varied with time

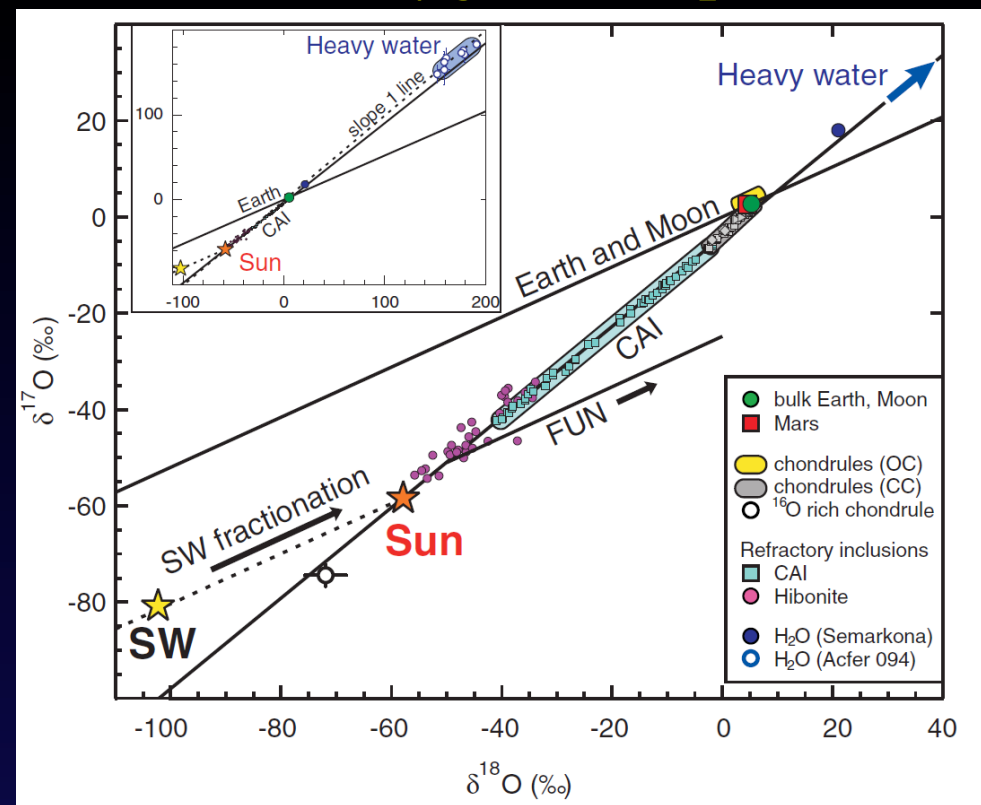
# Water as a carrier of heavy oxygen in the molecular cloud & PPD



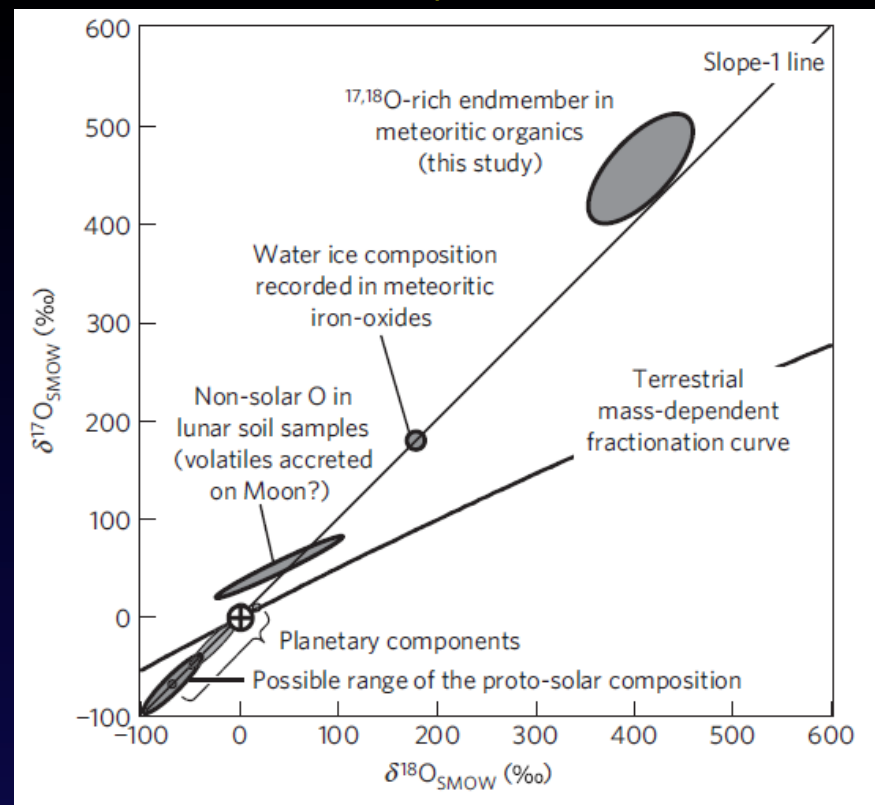
- preferential photodissociation of  $C^{17}O$  &  $C^{18}O$  in the **initially uniformly  $^{16}O$ -rich** ( $\Delta^{17}O \sim -25\text{‰}$ ) PPD or MC (**assumption**)
- released  $^{17}O$  &  $^{18}O$  incorporated into  $H_2O_{(s)}$ ;  $CO_{(g)}$  is  $^{16}O$ -enriched
- $H_2O_{(s)}/CO_{(g)}$  enrichment in the midplane of the protoplanetary disk followed by ice evaporation  $\rightarrow$   $^{17,18}O$ -enriched gas
- thermal processing of dust in  $^{17,18}O$ -rich gas  $\rightarrow$  evolution of solids towards the TFL



# Oxygen isotopes & water in the Solar System



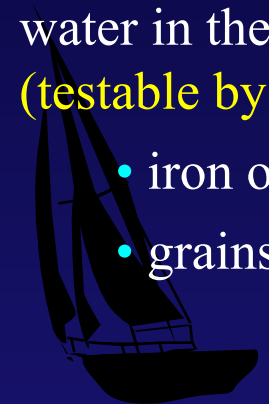
McKeegan et al. (2011) Science



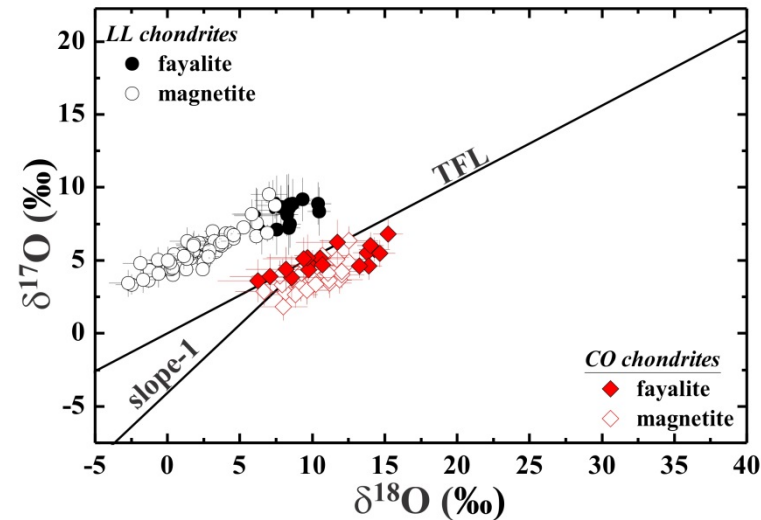
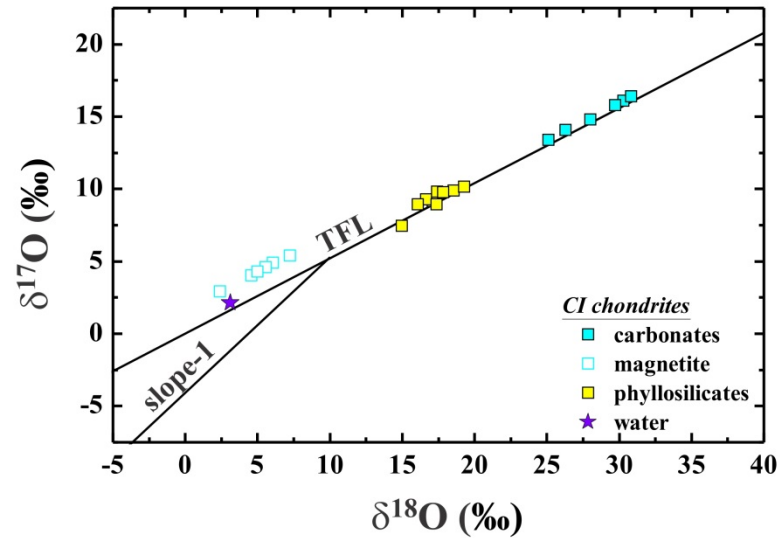
Hashizume et al. (2011) Nature Geoscience

- in the CO self-shielding models (Yurimoto & Kuramoto, 2004, Science; Lyons & Young, 2005, Nature), water in the outer disk (>30 AU) is highly  $^{17}\text{O}$  &  $^{18}\text{O}$ -enriched relative to the inner disk (testable by ALMA)

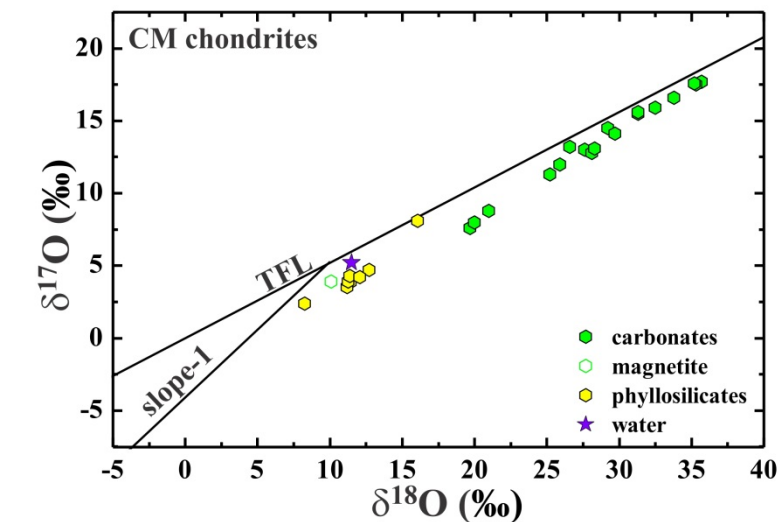
- iron oxides in Acfer 094 (ungr.):  $\Delta^{17}\text{O} \sim +90\text{‰}$  (Sakamoto et al., 2007, Science)
- grains in IOM from Y-793495 (CR):  $\Delta^{17}\text{O} \sim +500\text{‰}$  (Hashizume et al., 2007)



# Oxygen-isotope compositions of aqueously-formed minerals



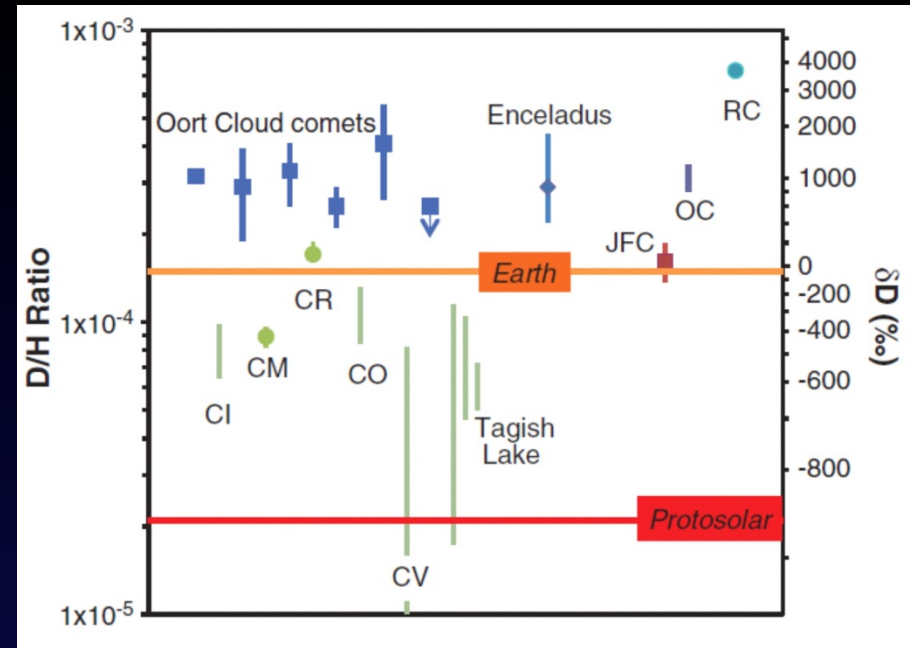
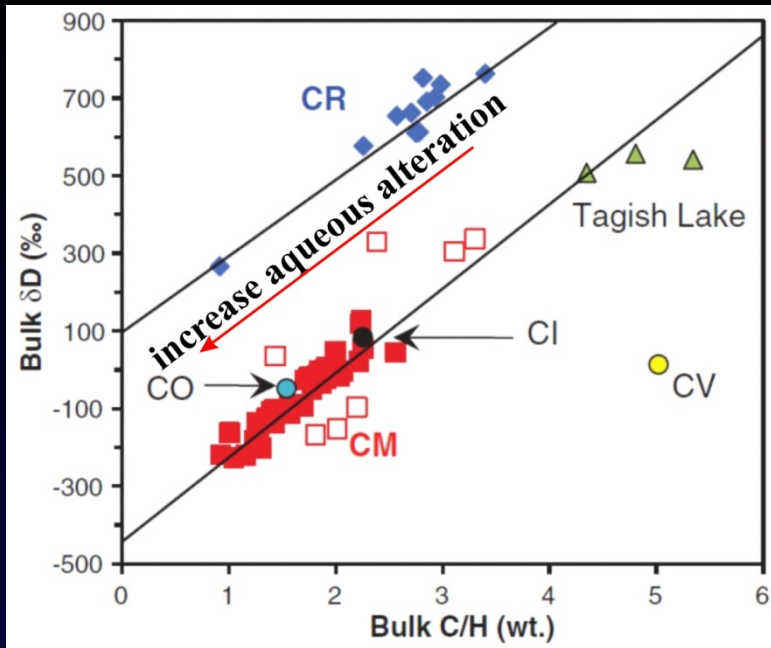
*Krot et al. (2012, 2013)*



- $\Delta^{17}\text{O}$  of aqueously-formed minerals can be used as a proxy for  $\Delta^{17}\text{O}$  of asteroid water ices; it stays  $\sim$ constant during alteration
- near terrestrial  $\Delta^{17}\text{O}$  values of water ices  $\rightarrow$  local, inner SS origin of water



# Sources of water on asteroids: Insights from D/H ratio

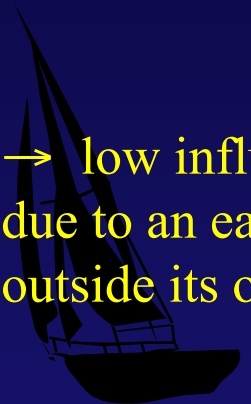


Alexander et al. (2012) Science

- $\delta D = (D/H_{\text{sample}}/D/H_{\text{SMOW}} - 1) \times 1000$
- bulk  $\delta D = \text{phyllosilicates} + \text{organics}$
- $\delta D$  of phyllosilicates at  $C/H = 0$

- $\delta D$  in chondrites, Oort Cloud & Jupiter Family comets, & Enceladus
- water in chondrites & comets formed in different SS regions (contrary to Walsh et al., 2011, Nature)

→ low influx of water from the outer SS into the inner disk  $\sim 2$  Myr after  $t_0$  could be due to an early growth of Jupiter that prevented significant radial transport of dust from outside its orbit



# Constraints on the early SS processes from chondrites

- ✓ Solar System formed near massive star(s)
- ✓ early generation of  $^{16}\text{O}$ -rich &  $^{16}\text{O}$ -poor reservoirs in protoplanetary disk
- ✓ inner Solar System solids experienced extensive thermal processing during evaporation, condensation, thermal annealing, & melting
- ✓ thermally processed solids were radially transported to the outer Solar System
- ✓ life-time of the PPD is  $\sim 3\text{-}4$  Myr
- ✓ accretion of cm-sized objects started at  $t_0$
- ✓ accretion of asteroid-sized bodies started  $< 1$  Myr after  $t_0$  & lasted at least 3-4 Myr
- ✓ accretion of individual asteroids may have been very rapid
- ✓ most chondrites accreted water ices, i.e., were close to the Snow Line
- ✓ influx of the outer Solar System material was small during chondrite accretion

