

Water in protoplanetary disks: D/H ratio

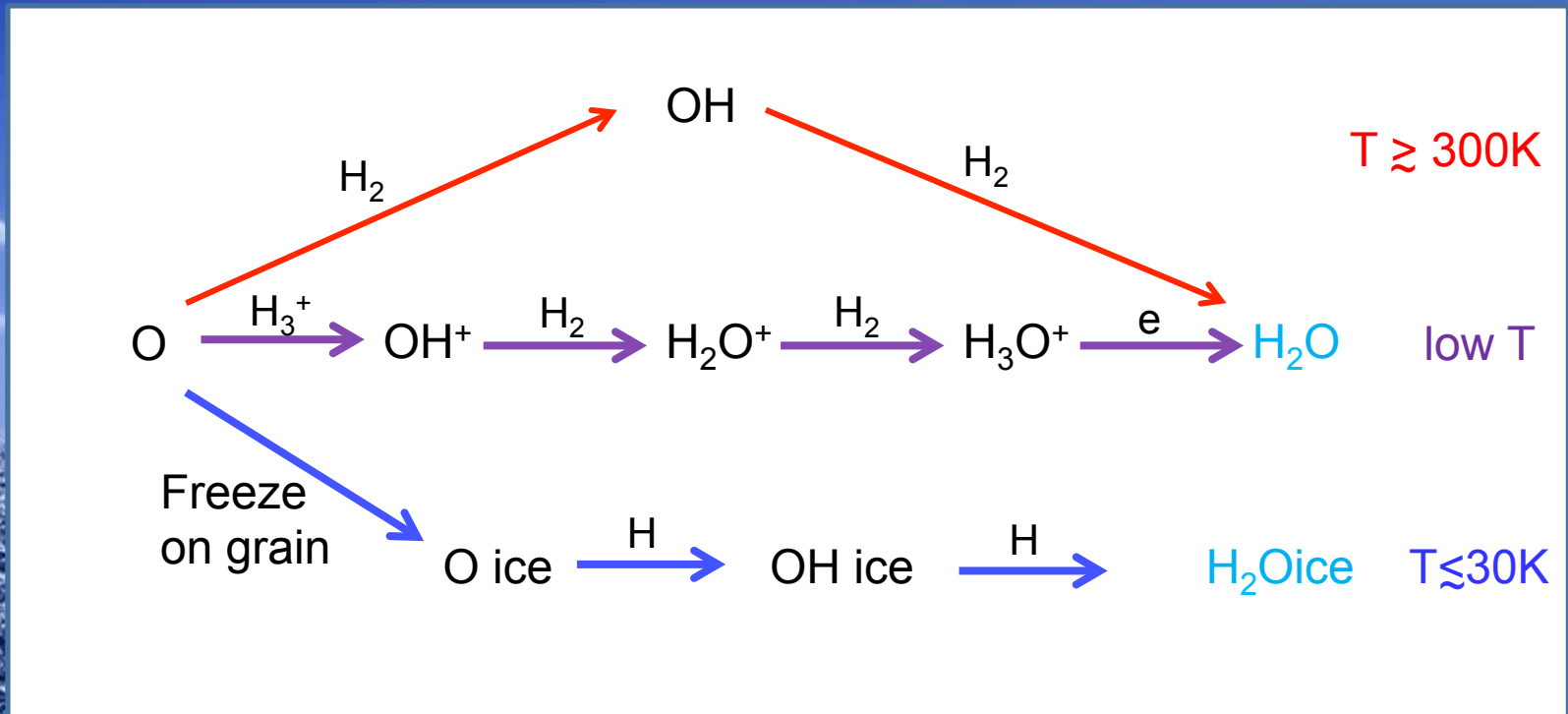
Kenji Furuya , Yuri Aikawa (Kobe Univ)

Hideko Nomura (Kyoto Univ)

Franck Hersant, Valentine Wakelam (Bordeaux Obs)

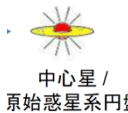
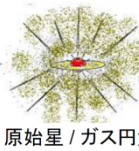
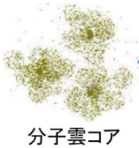
Water

- Major Oxygen reservoir
- Major ice (+ other volatiles as impurities)
 - ... raw material for {
 - outer planets
 - atmosphere
 - ocean
- 3 formation paths of water in ISM



Water in star-forming regions

Evolutionary stage	Water vapor $n(\text{H}_2\text{O})/n_{\text{H}}$	Water ice column density [cm^{-2}]
Starless core	$\sim 10^{-9}$	$\sim 10^{18} \text{ cm}^{-2}$ 10% of elemental O
Protostellar core (Class 0--I):		
Outer envelope ($T < 100 \text{ K}$)	$10^{-9} - 10^{-8}$	$\sim 10^{18} \text{ cm}^{-2}$
Inner envelope ($T > 100 \text{ K}$)	$10^{-6} - 10^{-5}$	---
Cavity wall by outflow	$10^{-5} - 10^{-4}$	---
Circumstellar disk (Class II):		
Outer disk ($> 100 \text{ AU}$)	$< 10^{-7}$	$\sim 10^{18} \text{ cm}^{-2}$
Inner disk ($< 3 \text{ AU}$)	1.3 [relative to CO]	---



Ref. Whitet & Duley (1991); Terada et al. (2007); Najita & Carr (2008); Honda et al. (2009);
Kristensen et al. (2010, 2012); Hogerheijde et al. (2011); Caselli et al. (2012); Coutens et al. (2012)

Water is everywhere!

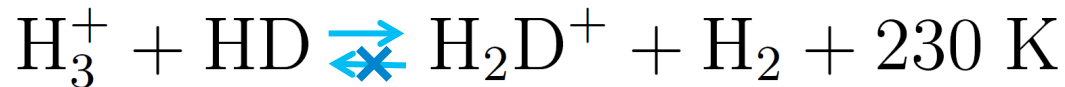
→ Does water in disks & planetary systems originate in ISM?

HDO/H₂O ratio

- Probe of formation site of H₂O

formation at lower temperature → Higher HDO/H₂O ratio

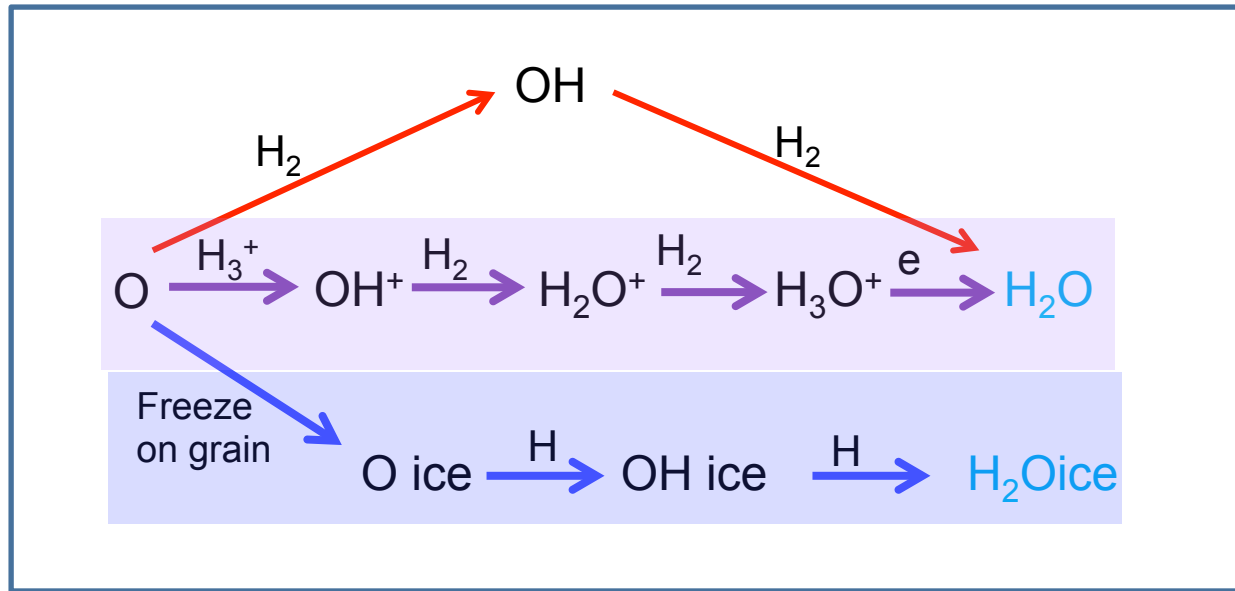
Exothermic Exchange Reactions



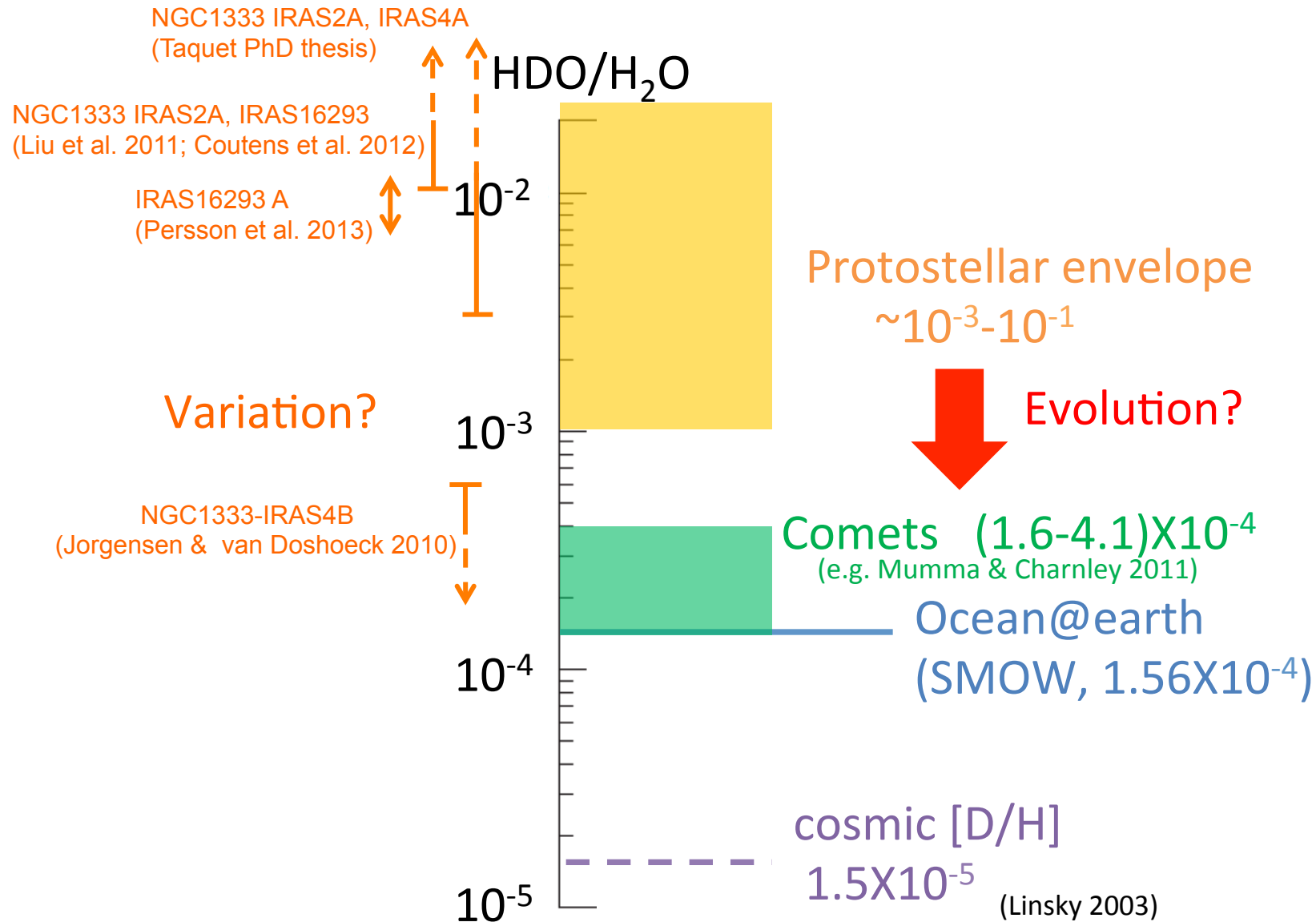
- high H₂D⁺/H₃⁺

(e.g. Millar, Bennet & Herbst 1989)

- H₂D⁺ + e → H₂ + D → high atomic D/H

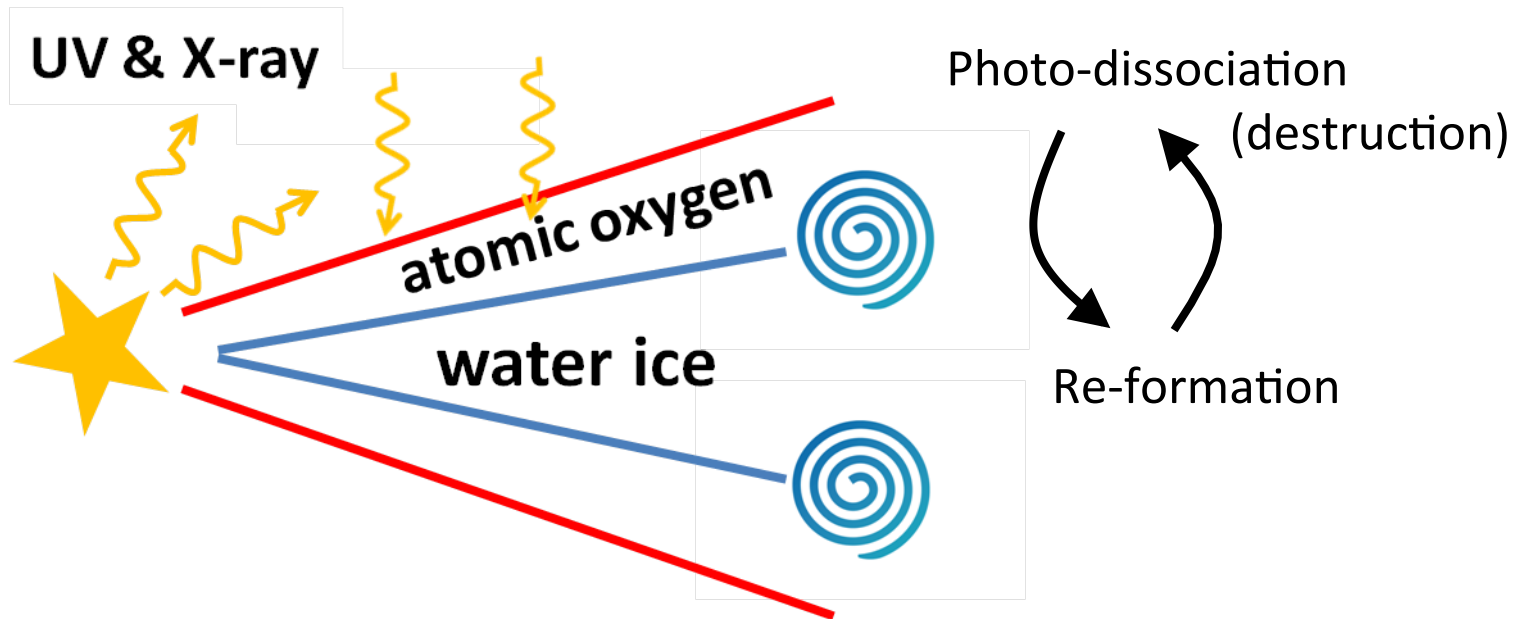


HDO/H₂O in comets and protostellar cores



This Work

- Can water be destroyed and re-formed in disks?
 - Does HDO/H₂O ratio change?
- Effect of vertical mixing by turbulence



(cf. Lyons & Young 2005)

Model

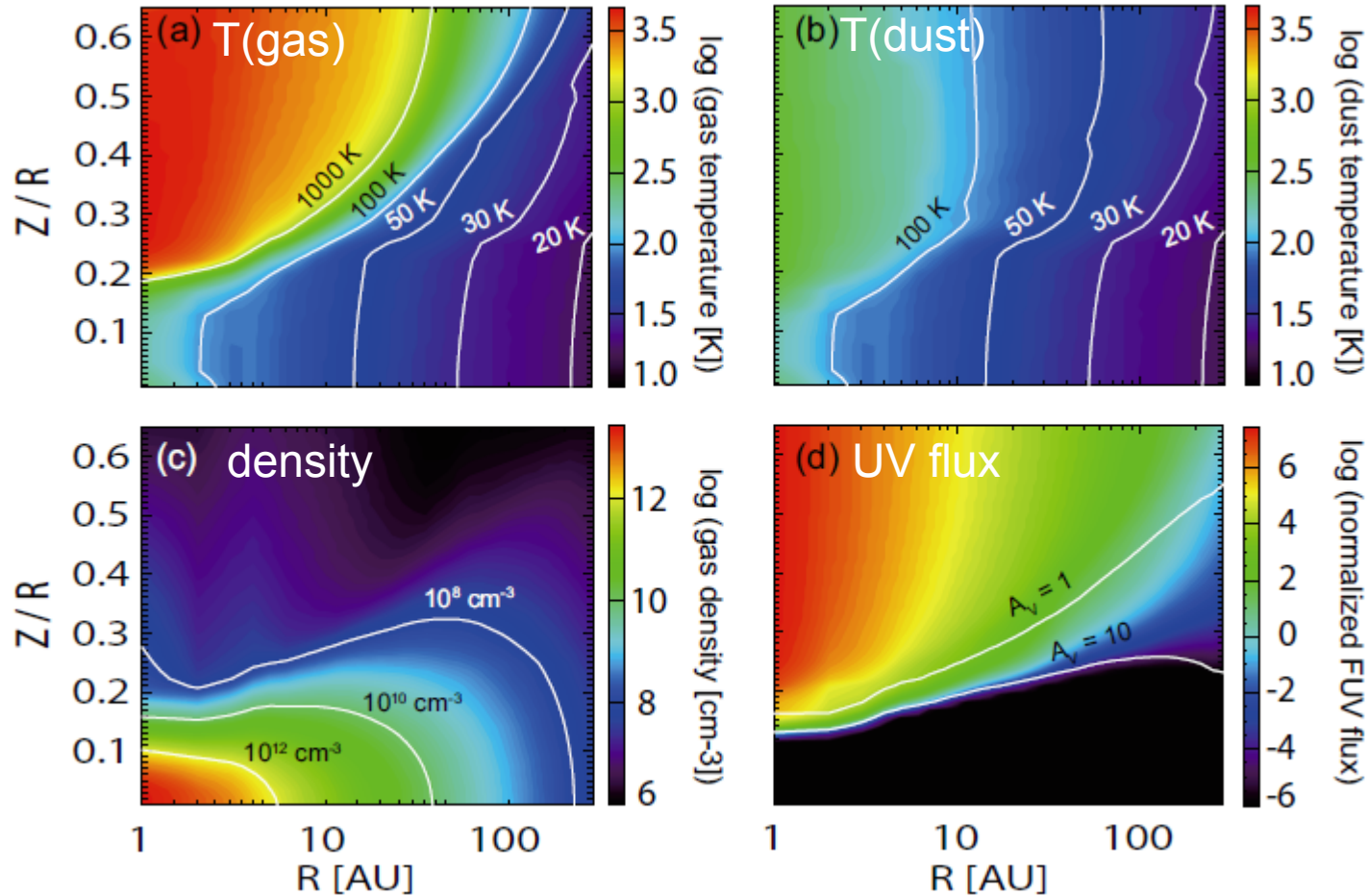
➤ Disk model (Nomura et al. 2007)

- Small dust (i.e. large surface area) well-mixed with gas
- Stellar, interstellar, and induced UV
- Stellar X-ray

$$M_* = 0.5 M_\odot$$

$$L_* = 0.87 L_\odot$$

$$M_d = 2.4 \times 10^{-2} M_\odot$$



Model

➤ Chemical model

Gas – Grain network of Garrod & Herbst (2006)

+ mono- & multi-**Deuteration** (Aikawa et al. 2012; Talukdar et al. 1996; Bergin et al. 1999; UMIST07)

+ **high T reactions** (Harada et al. 2010)

Photodesorption & photodissociation of ice

(Oberg et al. 2007; Andersson & van Dishoeck 2008; Mason et al. 2006)

- Desorption yield $\sim 10^{-3}$

- Dissociation is allowed only in the surface 2 layers

➤ Basic equation (e.g. Xie et al. 1995; Willacy et al. 2006)

$$\frac{\partial n_i}{\partial t} - \underbrace{\frac{\partial}{\partial z} \left[n D_z \frac{\partial}{\partial z} \left(\frac{n_i}{n} \right) \right]}_{\text{diffusion(vertical)}} = \underbrace{P_i - L_i}_{\text{Chemistry}}$$

n_i : num density of species i

grain radius: $0.1 \mu\text{m}$

$D_z = \alpha_z c_s h$: diff coef

$\alpha_z = 0, 10^{-3}, 10^{-2}$

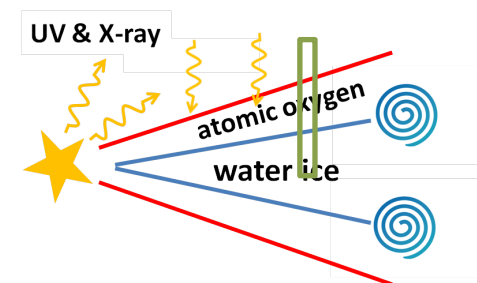
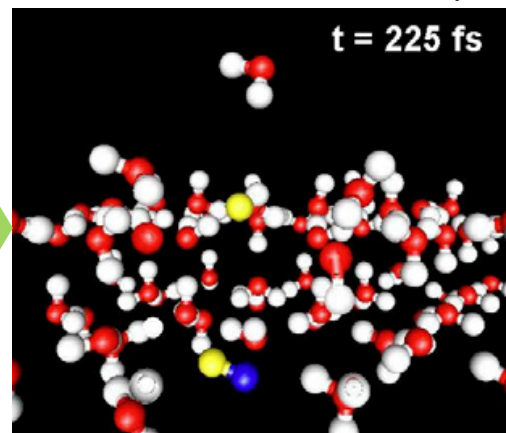
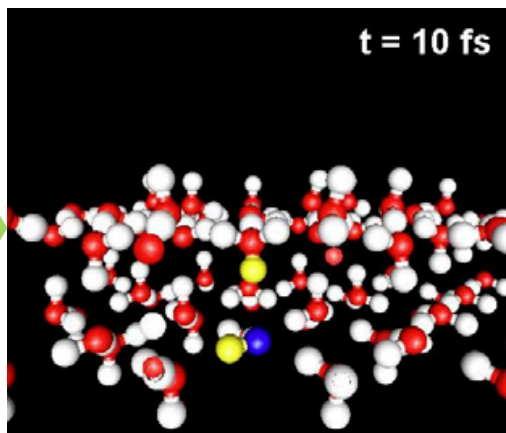
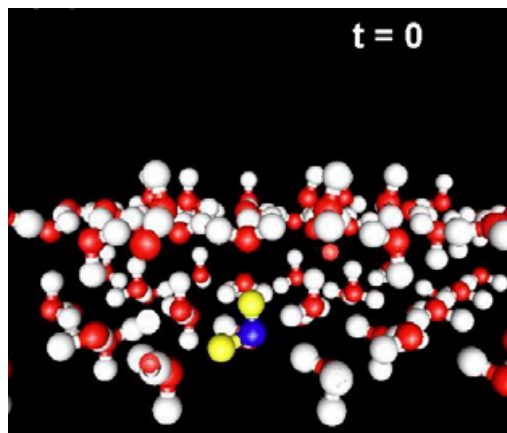
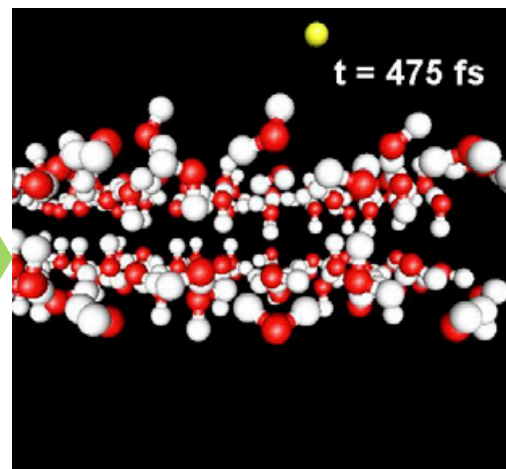
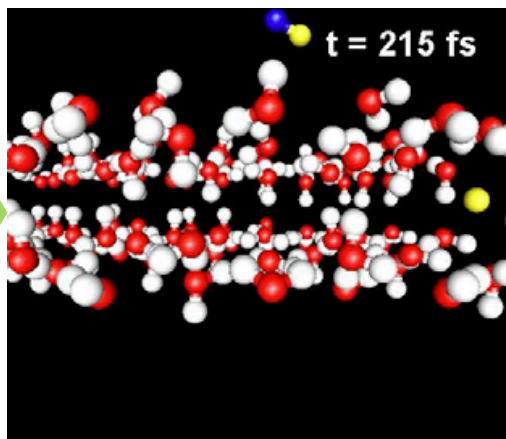
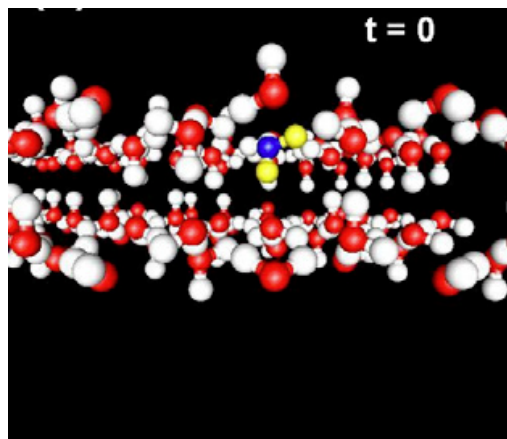


Photo-desorption



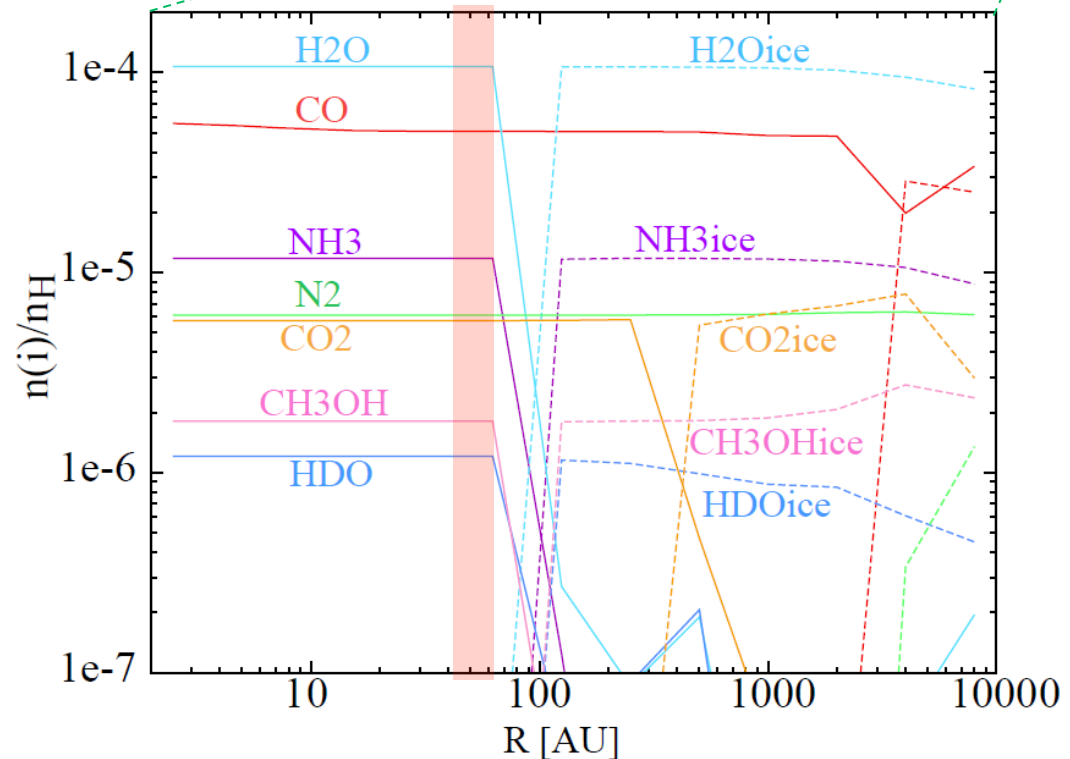
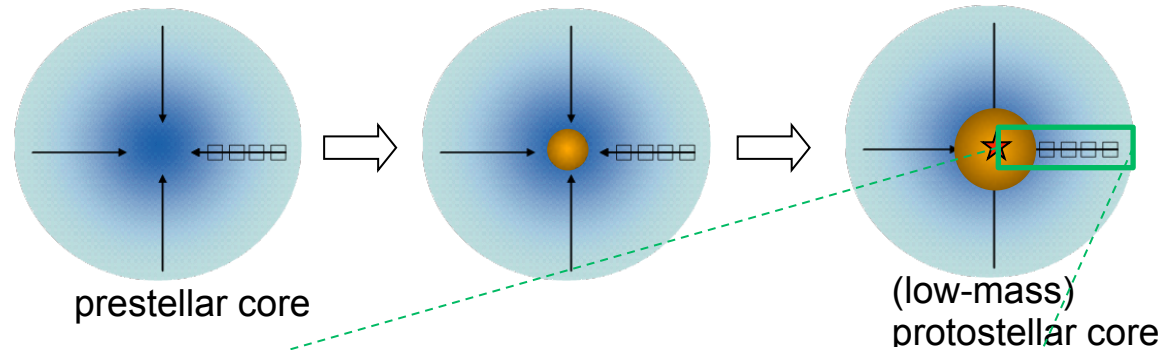
Andersson & van Dishoeck (2008)

Photo-dissociation of ice



➤ Initial Abundances

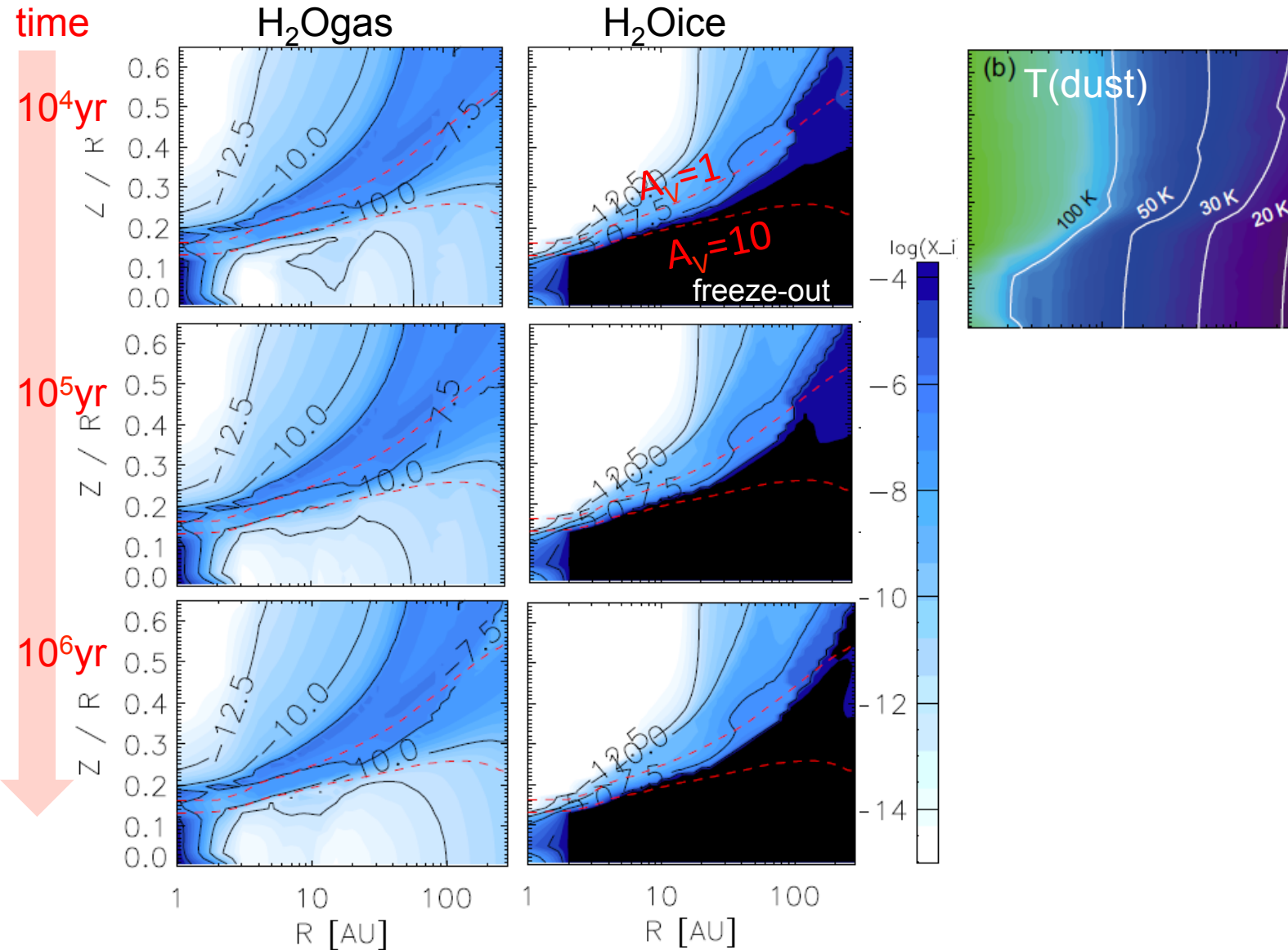
- Collapsing star-formation core (Aikawa et al. 2012)
- H₂O is the major O carrier
- HDO/H₂O ~ 10⁻²



@ 9.3×10^4 yr

Result: H₂O abundances

$\alpha=0$: no diffusion



Result: H₂O abundances

$\alpha = 10^{-2}$

time

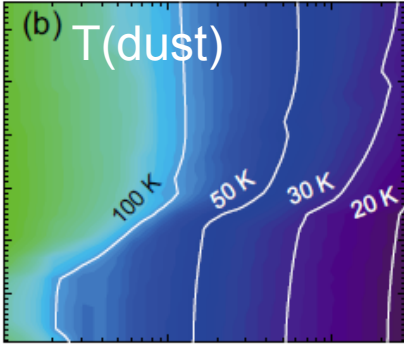
10⁴yr

10⁵yr

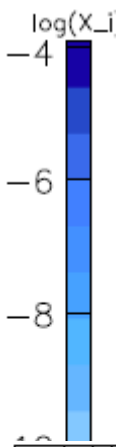
10⁶yr

H₂Ogas

H₂Oice

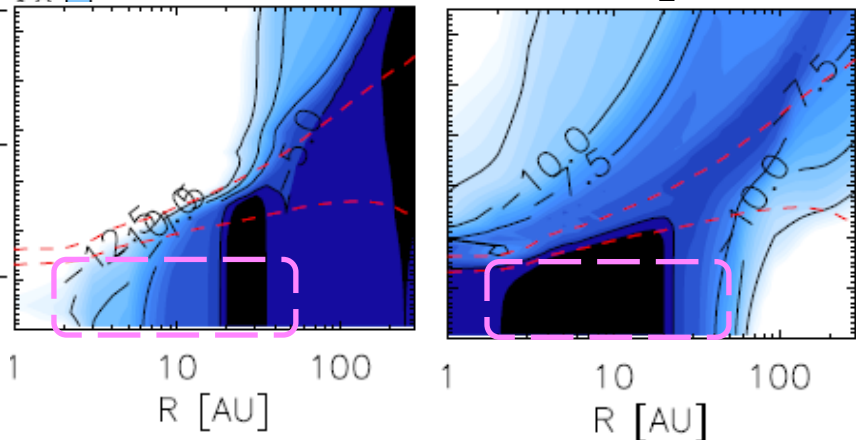
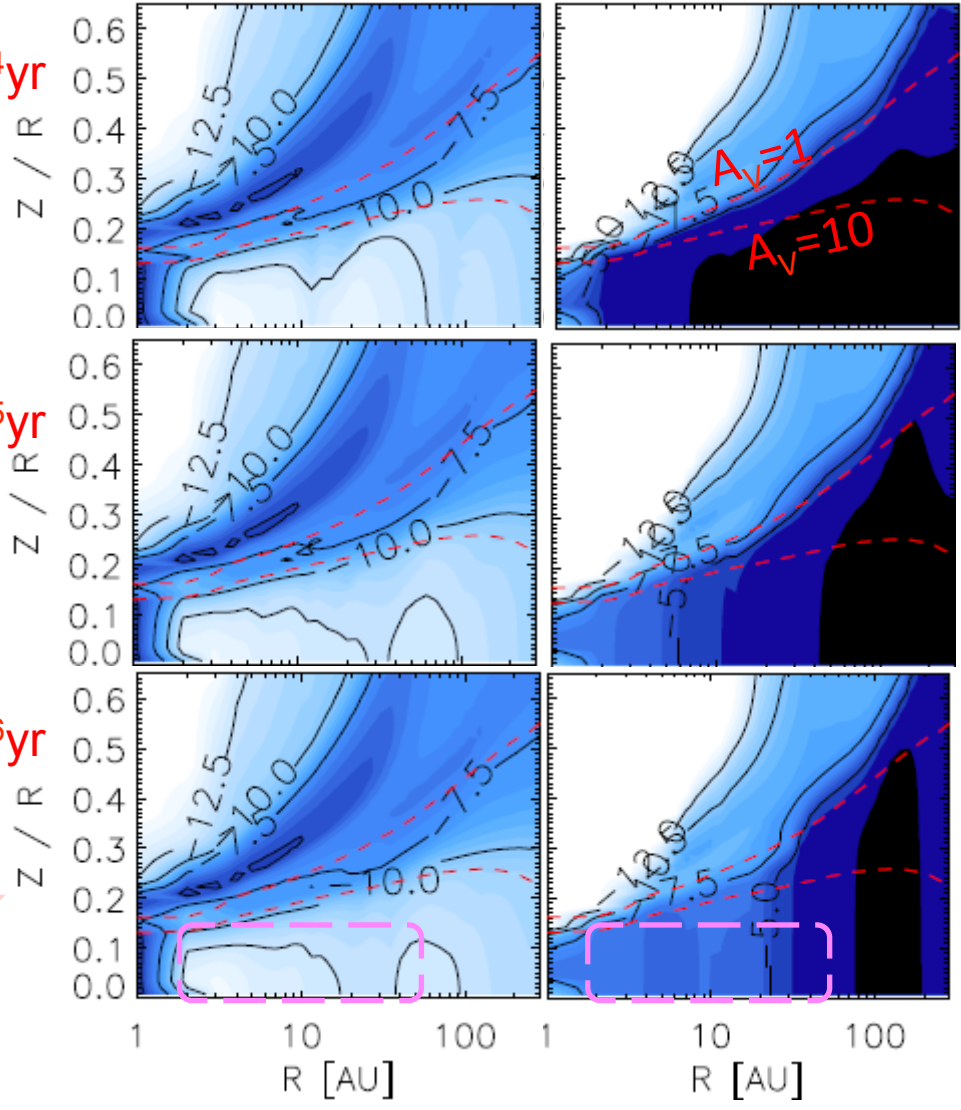


-Water is gone at $r < 40$ AU!
Where is Oxygen??



CO₂ice

O₂gas



Result: Oxygen chemistry

- H_2O_{ice}

Transported to disk surface
 → photodissociation

- **O atom**

Transported to deeper layers

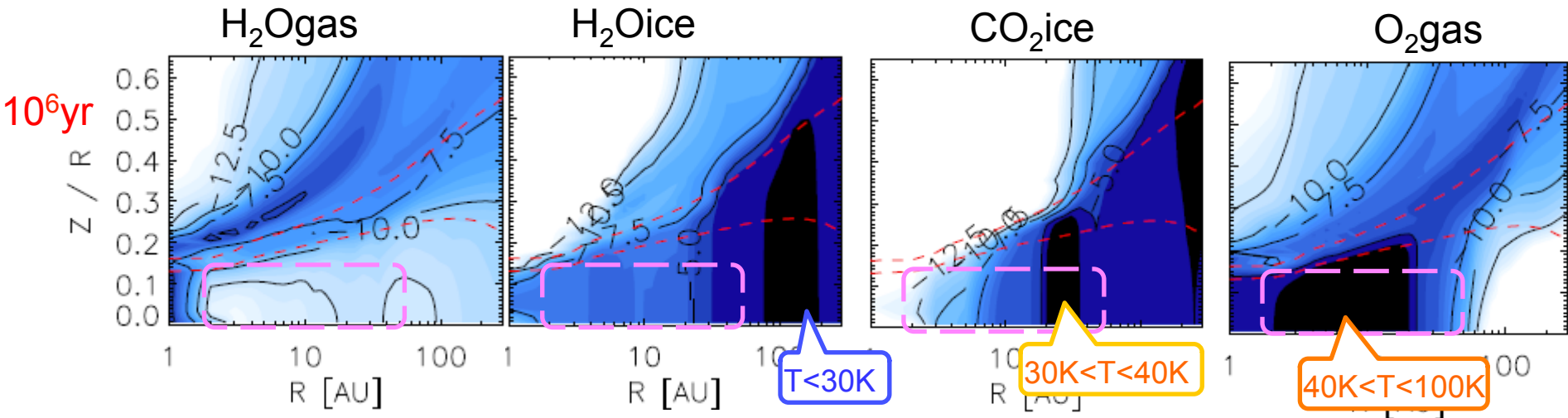
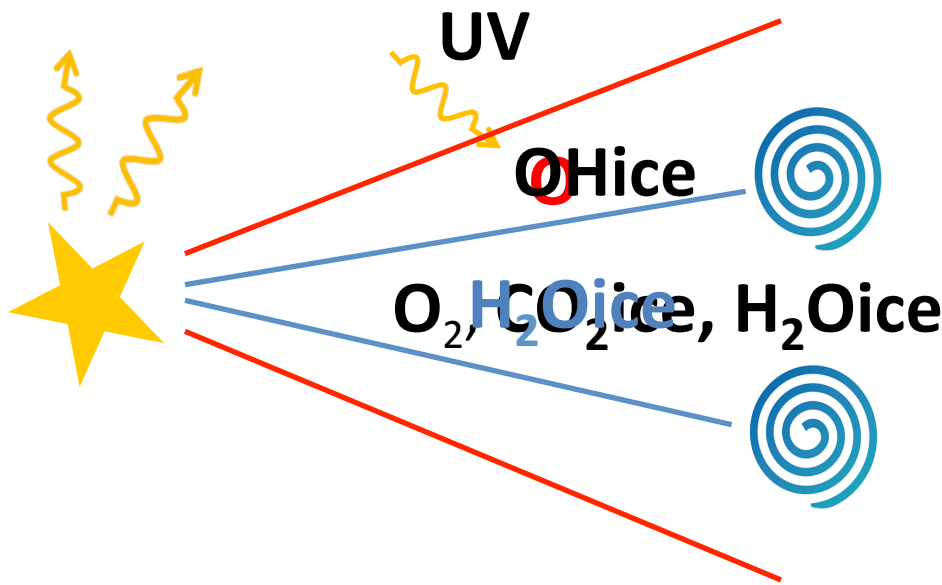
@ $T_d < 30K$

H_2O ice is re-formed

@ $T_d > 30K$

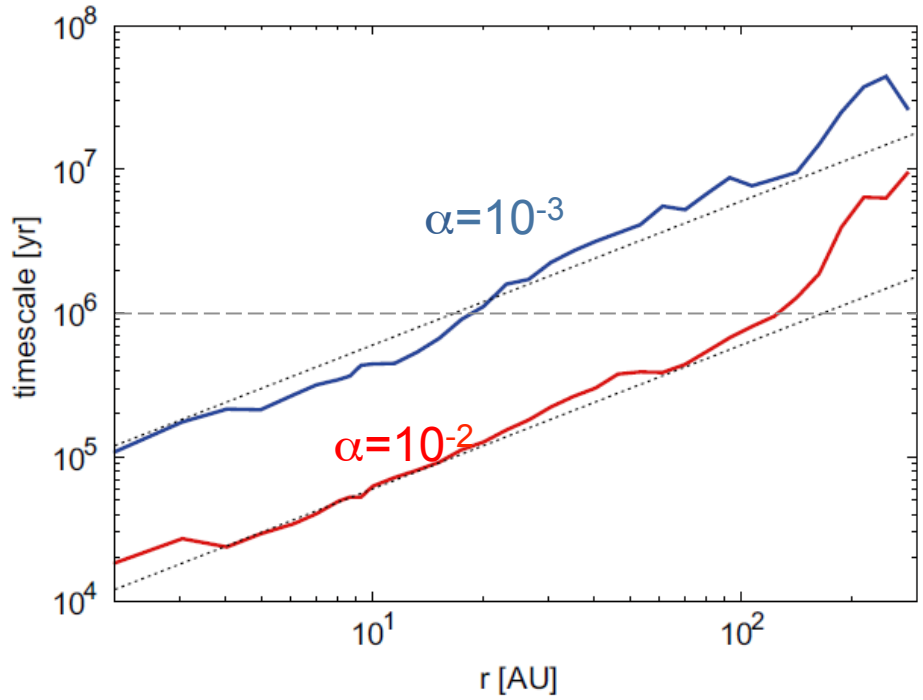
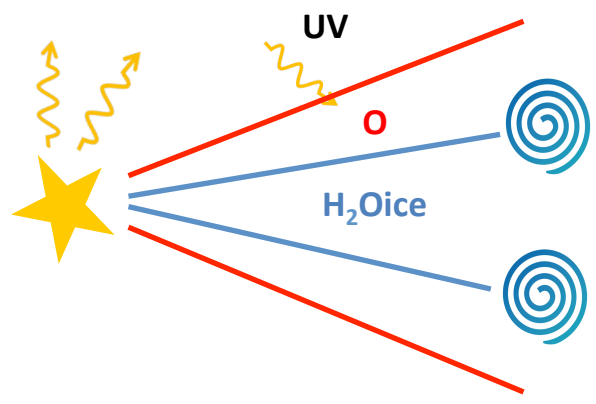
H_2O formation is inefficient

→ O is converted to O_2 & CO_2_{ice}

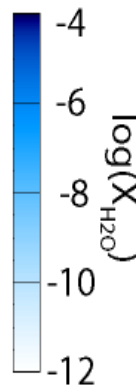
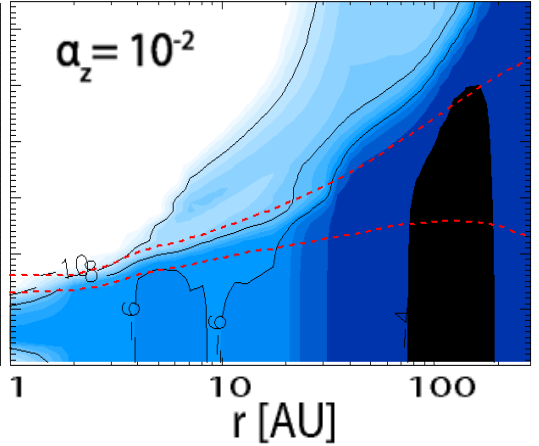
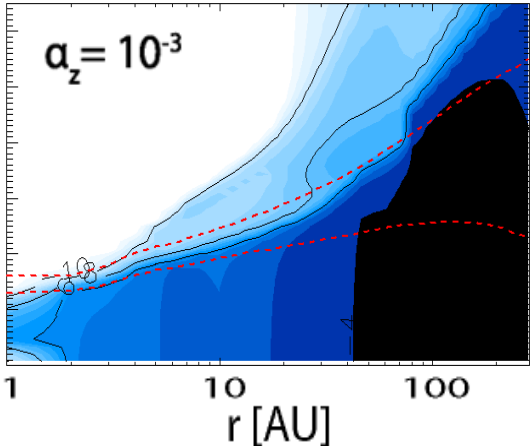
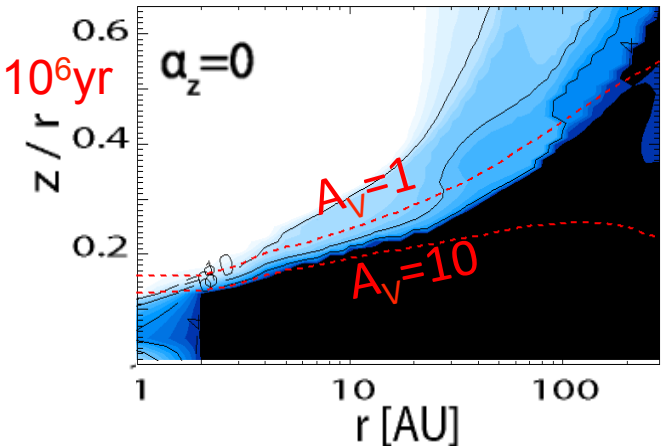


Result: Timescale of H₂O destruction

$$\tau \approx \frac{N_{\text{H}_2\text{Oice}}}{\text{flux @ H}_2\text{O - Oatom boundary}} \propto r$$



alpha ➔



Result: HDO/H₂O ratio

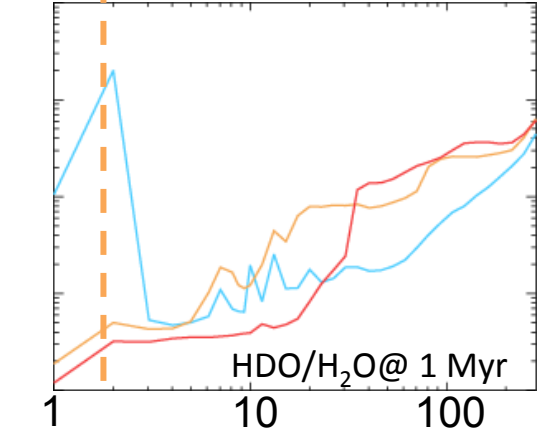
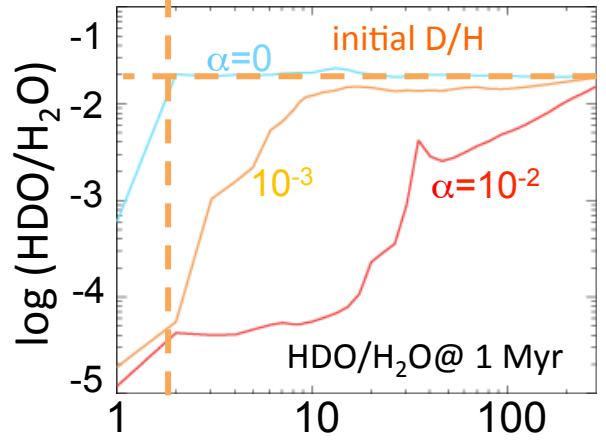
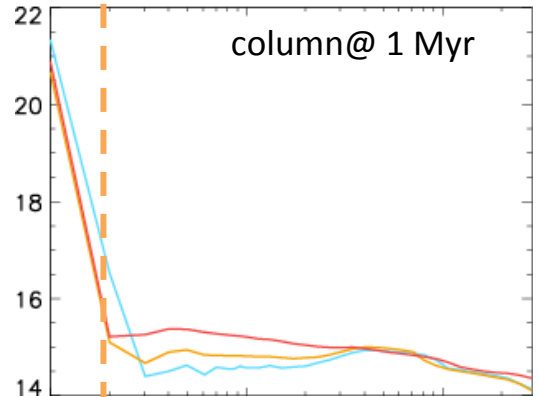
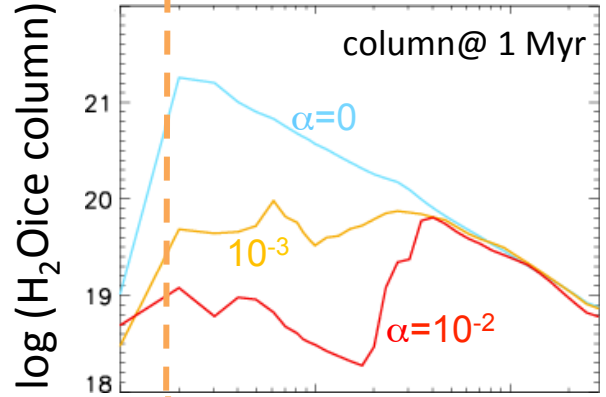
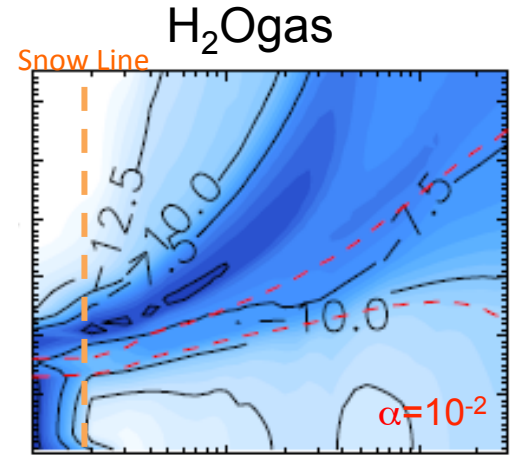
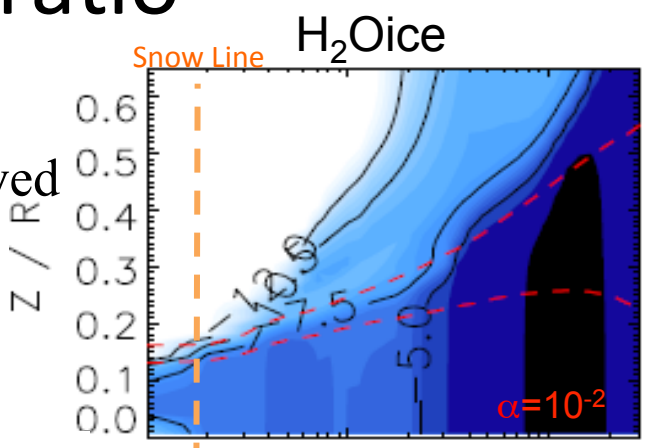
- HDO/H₂O ice ratio decreases ONLY when H₂O ice is destroyed significantly

- H₂O gas is abundant at
 { disk surface
 inside snow line

$$\left(\frac{\text{HDO}}{\text{H}_2\text{O}}\right)_{\text{gas}} < \left(\frac{\text{HDO}}{\text{H}_2\text{O}}\right)_{\text{ice}}$$

← Reformation in gas phase

(see also Willacy & Woods 2009)



Discussion: Dependence on $E_{\text{des}}(\text{H})$

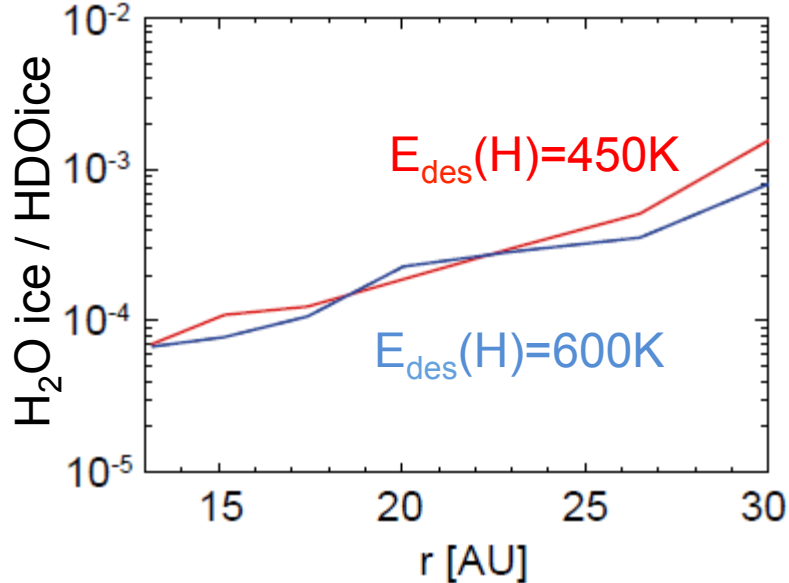
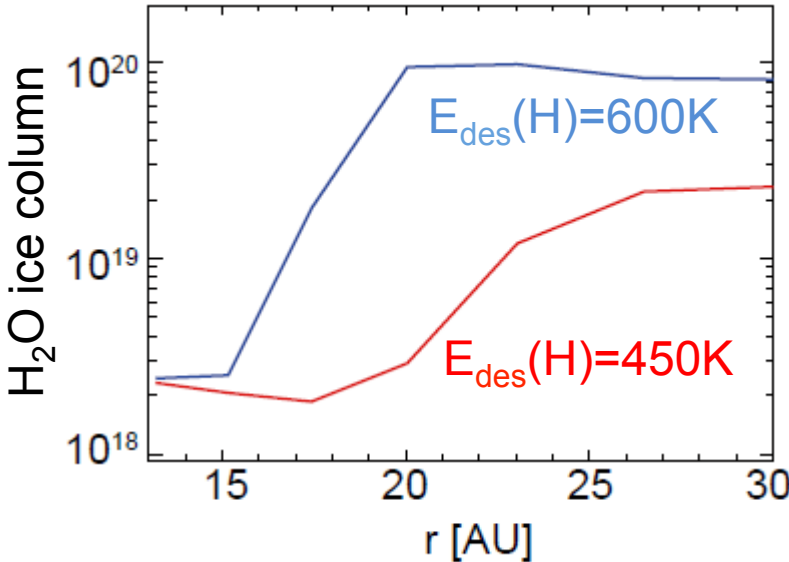
- H_2O reformation is inefficient, because H atom is easily desorbed
So far, we assumed $E_{\text{des}}(\text{H})=450\text{K}$



- $E_{\text{des}}(\text{H}) \sim 600\text{K}$ on amorphous ice
(Al-Halabi & van Dishoeck 2007; Watanabe et al 2010; Hama et al. 2012)

→ H_2O can be reformed at higher temperature!

→ Abundant H_2O ice with low D/H ratio $\sim 10^{-4}$



Summary on H₂O & Implications

- H₂O ice is photo-desorbed & dissociated at disk surface and
 - reformed at $T < \del{30K} 40K$
 - HDO/H₂O $\sim \del{10^{-3}-10^{-2}} 10^{-4}-10^{-3}$ if $E_{\text{des}}(\text{H}) = 600\text{K}$
 - converted to CO₂ at $30\text{K} < T < 40\text{K}$
 - converted to O₂ at $40\text{K} < T < 100\text{K}$
- Sedimentation of large grains (while small grains are stirred up) will protect H₂O ice. (cf. Hogerheijde et al. 2011)
- But then $(\text{HDO}/\text{H}_2\text{O})_{\text{ice}}$ does not change

- HDO/H₂O $\sim 10^{-4}$ in comets could be set in turbulent disk

- $(\text{HDO}/\text{H}_2\text{O})_{\text{gas}} \sim 10^{-4}-10^{-3}$ $\left(\frac{\text{HDO}}{\text{H}_2\text{O}}\right)_{\text{gas}} < \left(\frac{\text{HDO}}{\text{H}_2\text{O}}\right)_{\text{ice}}$

cf. $(\text{HDO}/\text{H}_2\text{O})_{\text{gas}} = 6 \times 10^{-4}$ @ NGC1333 IRAS4B (Jorgensen & van Dishoeck 2010)